

The role of ambient horizontal vorticity in near-ground rotation of supercells

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Acknowledgments:

George Bryan

Paul Markowski

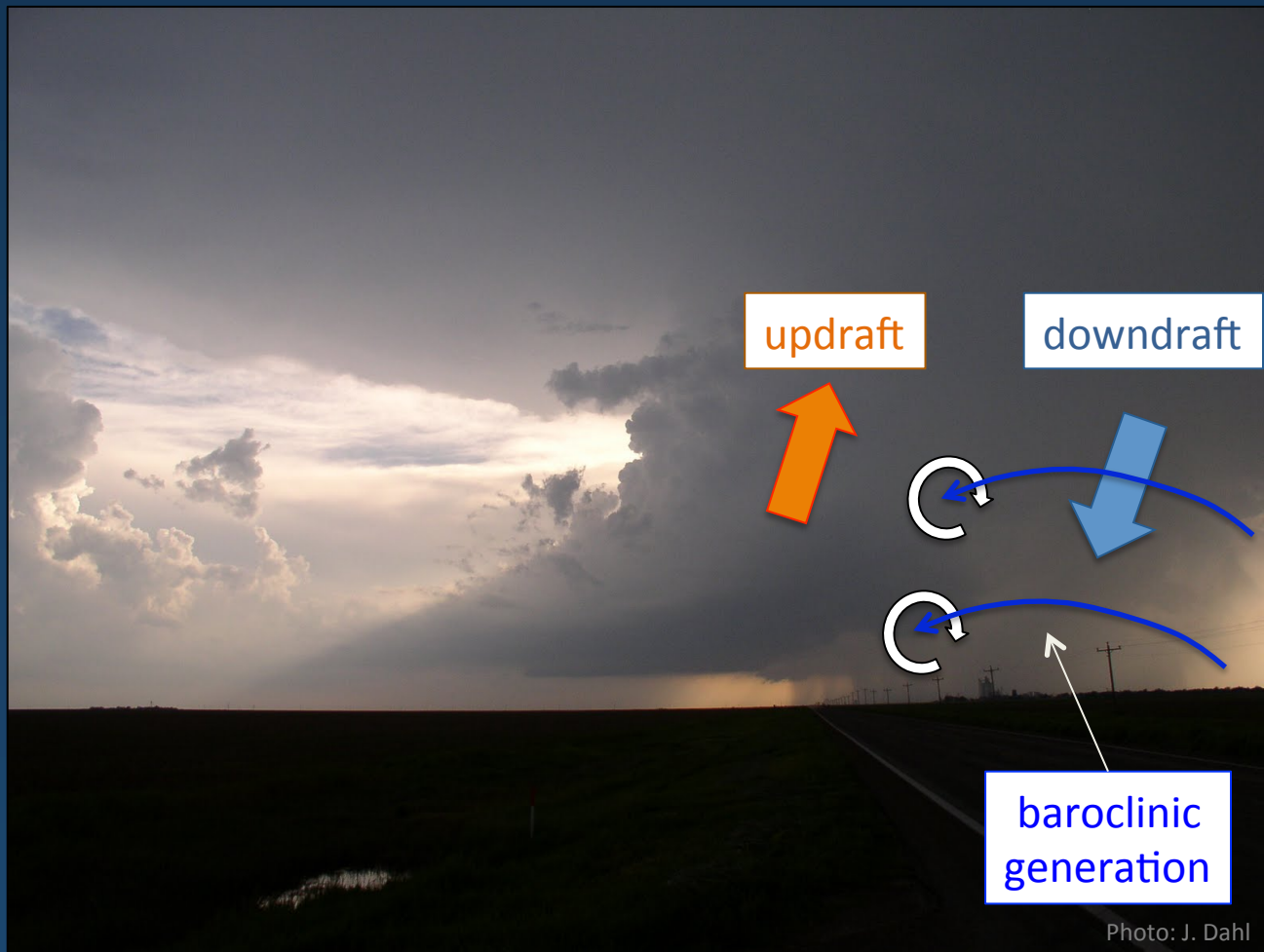
Bob Davies-Jones

Funding: National Science Foundation

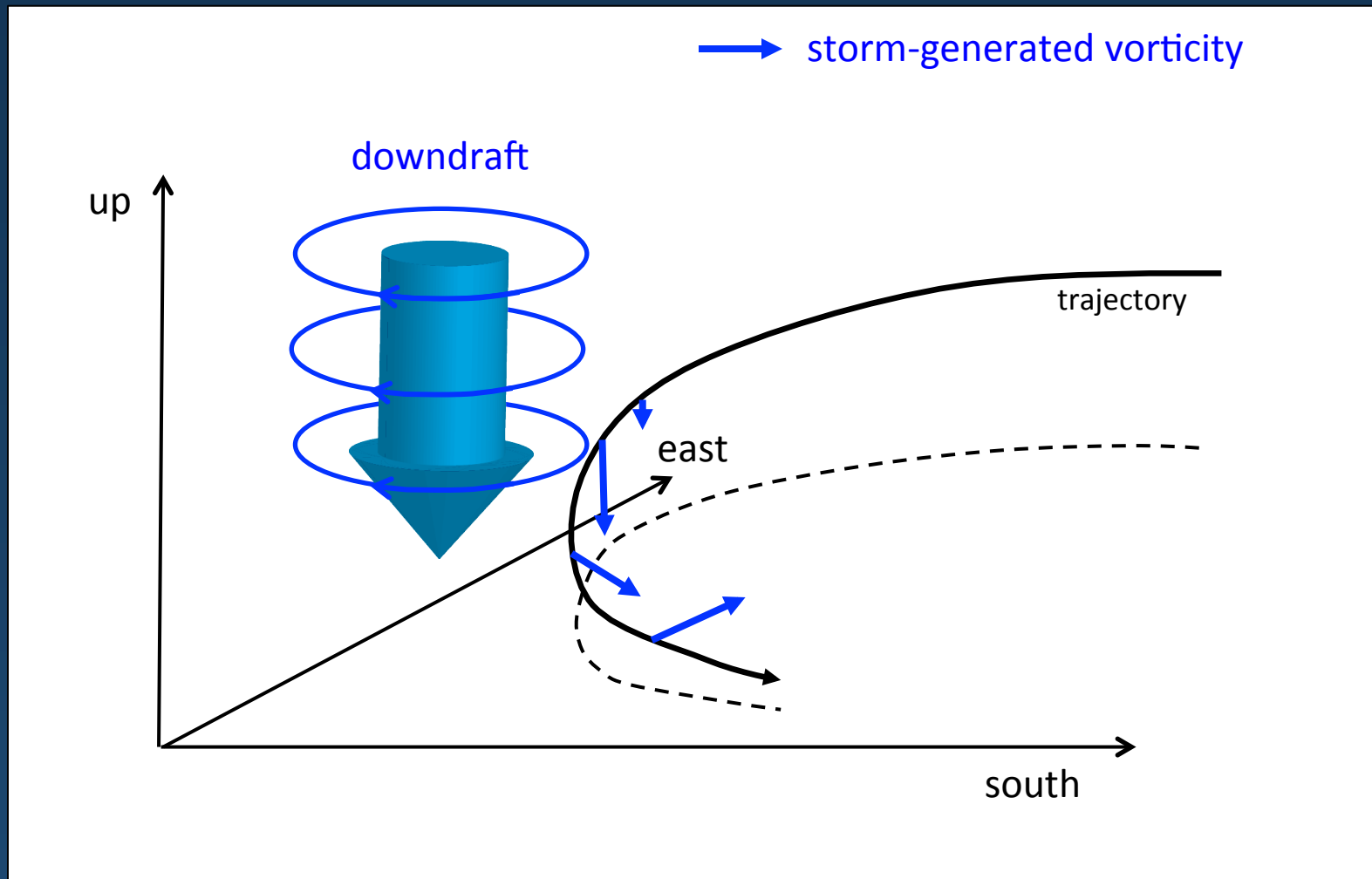
Background

- Supercell tornadogenesis proceeds in three steps:
 1. Rotation aloft (“midlevel mesocyclone”)
 2. **Rotation at the ground (requires downdraft)**
 3. Amplification of the surface rotation by convergence
- Numerous studies have demonstrated the importance of storm-generated (baroclinic) vorticity in surface rotation
- The role of *ambient* vorticity in near-surface rotation is not well understood

Baroclinic generation of vorticity

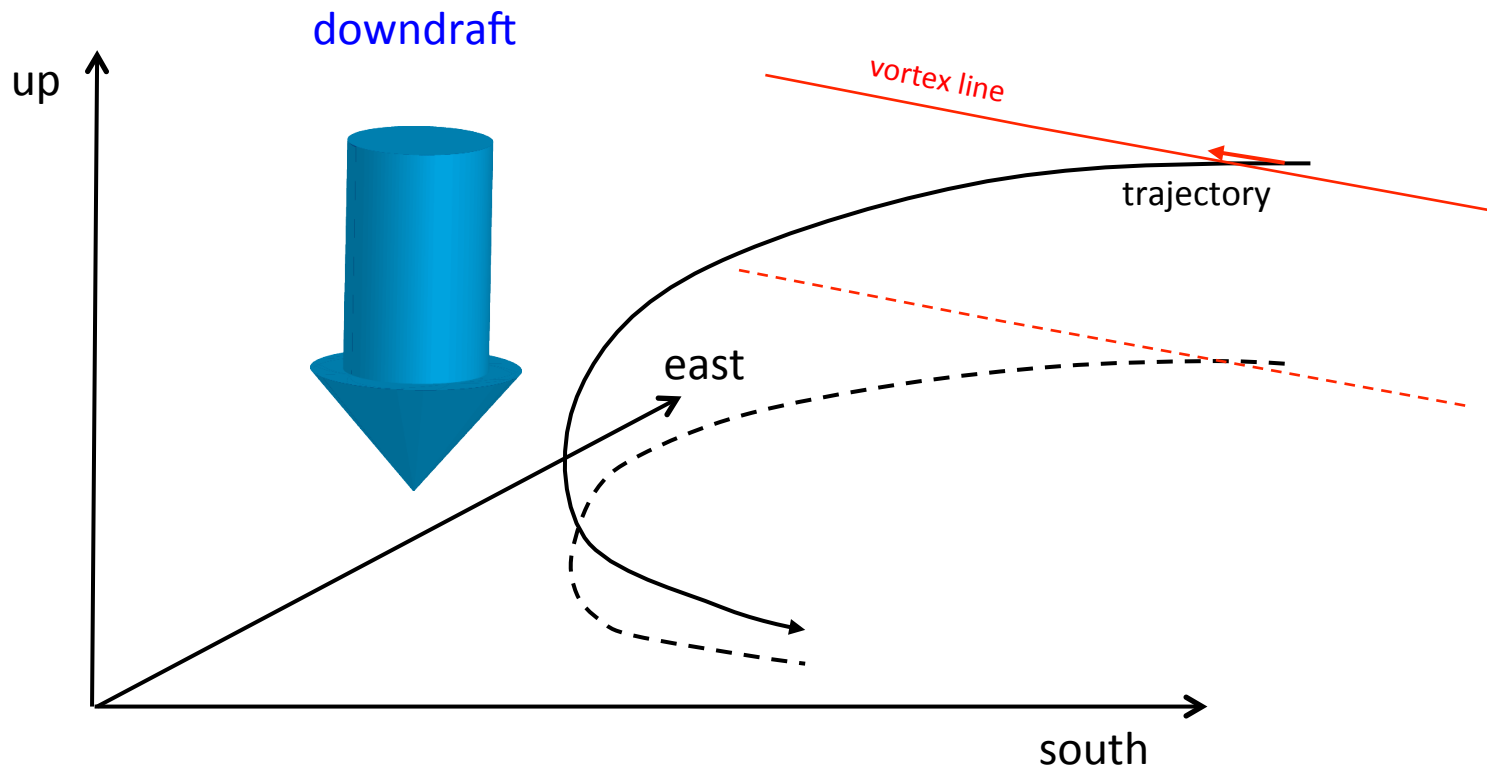


Baroclinic process (storm-generated vorticity)



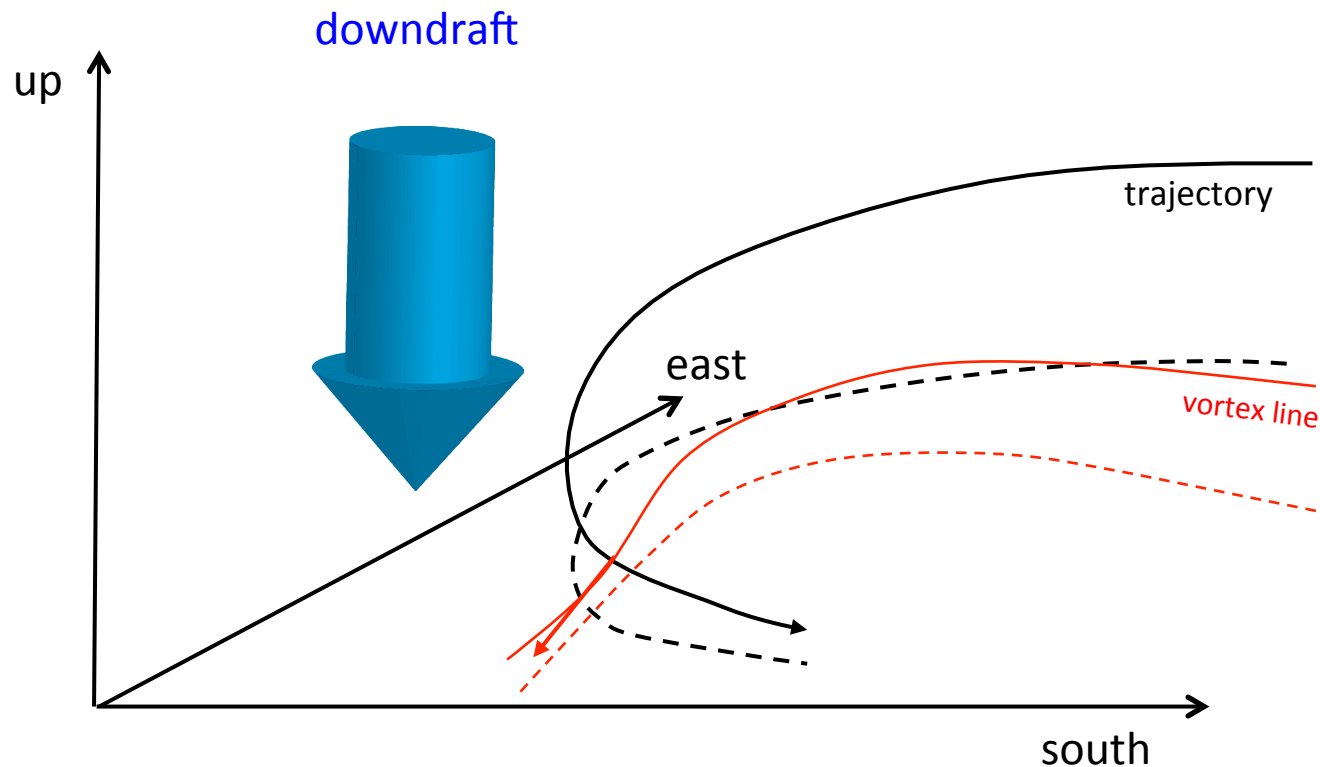
Barotropic process (imported vorticity)

Barotropic vorticity: Frozen into the fluid (rearrangement of *existing* vorticity)



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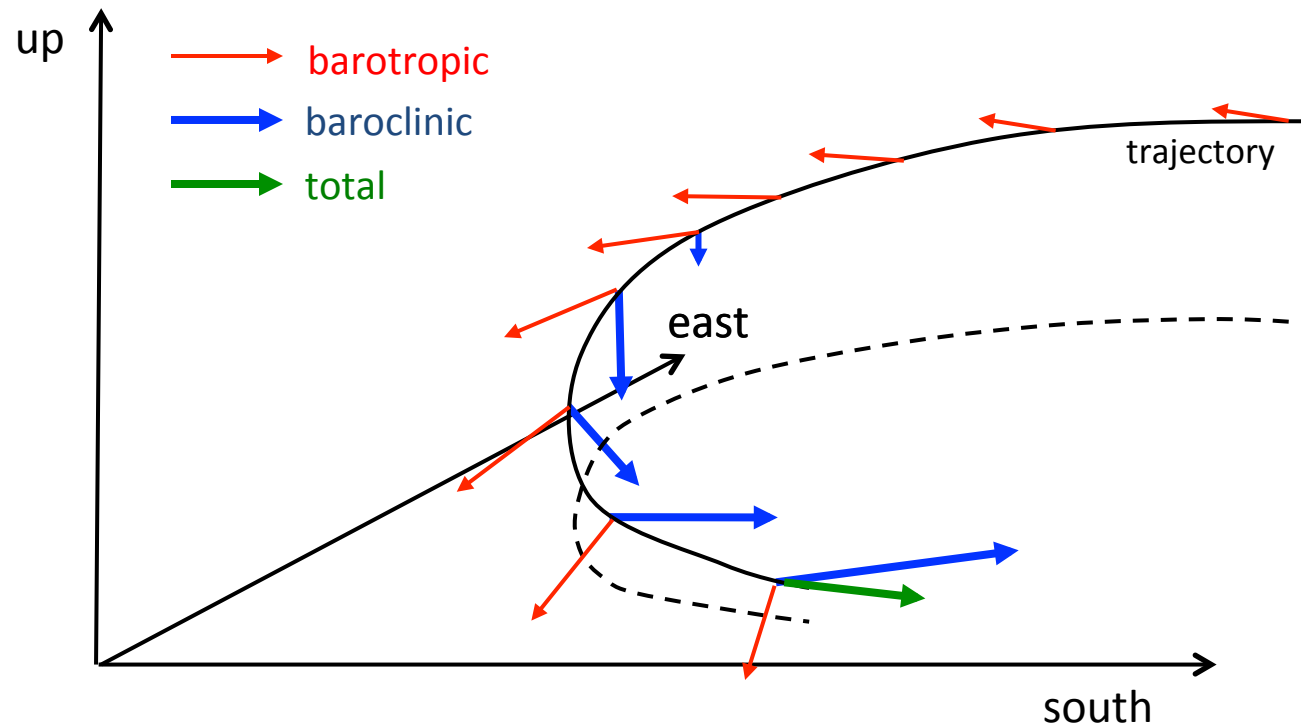
How does this barotropic vorticity contribute to the surface vorticity?

Vorticity decomposition

Barotropic vorticity: Frozen into the fluid (rearrangement of *existing* vorticity)

Baroclinic vorticity: Generated baroclinically and subsequently rearranged

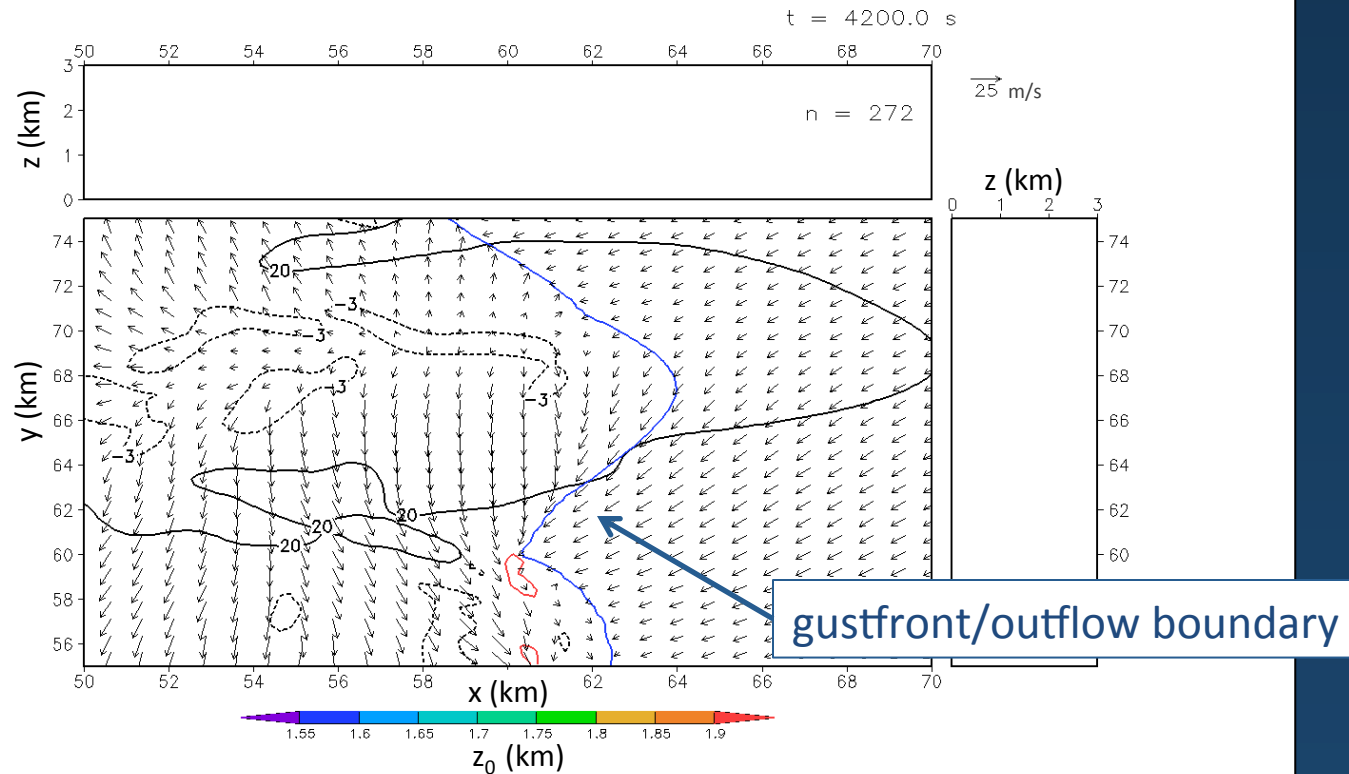
Total vorticity: Sum of baroclinic and barotropic parts



Simulation setup

- Bryan Cloud Model 1 (CM1), release 17
- $dx = dy = 250$ m
- $50 \text{ m} < dz < 250 \text{ m}$
- Single-moment microphysics
 - adjusted rain-intercept parameter to reduce cold-pool strength
- free-slip lower boundary
- Del City base state
- Initialization: 2K ellipsoidal potential-temperature perturbation
- Restart run with highly accurate forward trajectories
 - Calculated at each time step ($dt = 2$ s)
 - 4th-order forward integration
 - 4th-order spatial interpolation

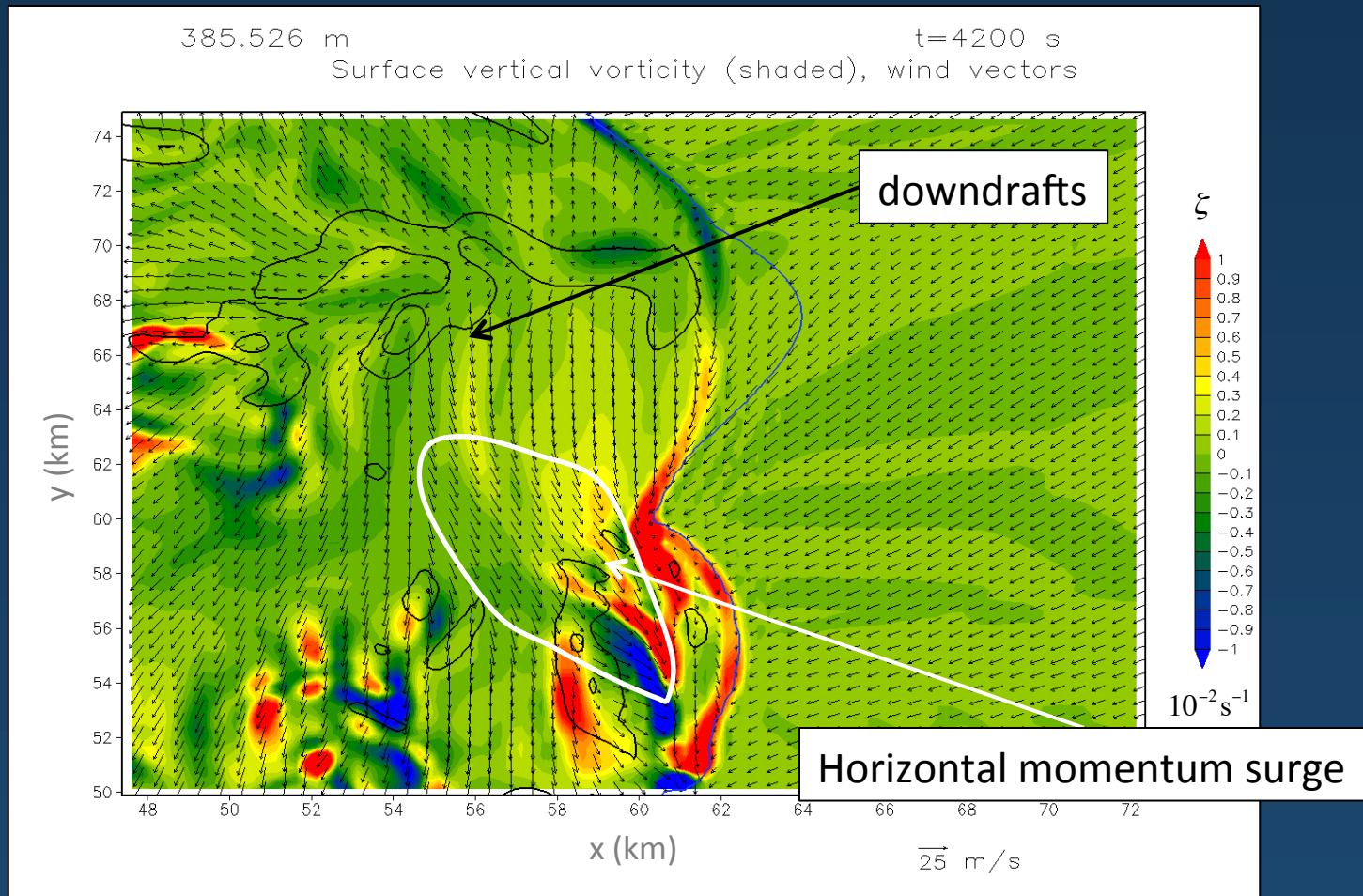
Forward trajectories



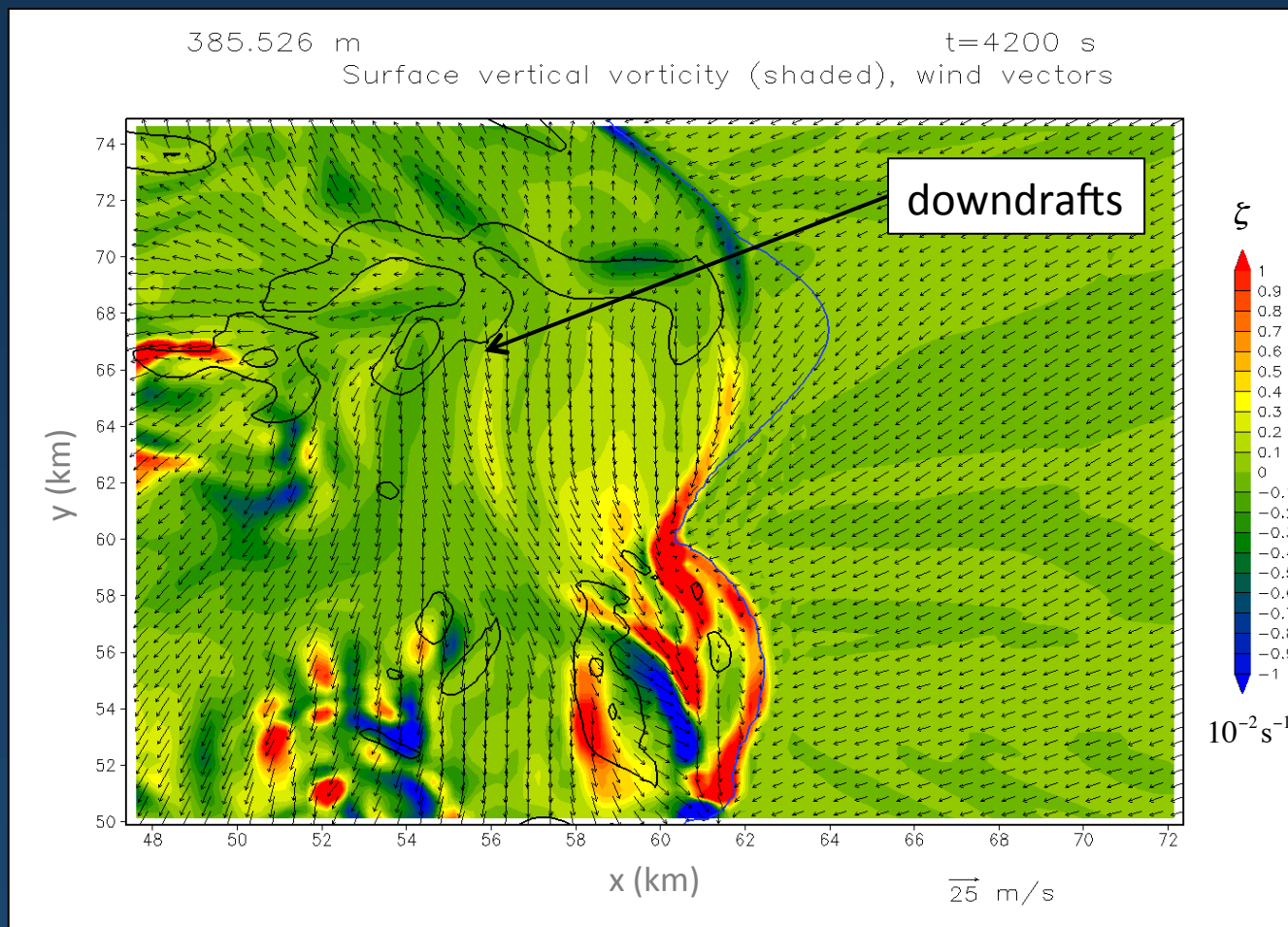
Black solid contour: reflectivity (385 m)
Black dashed contours: downdraft (385 m)
Red contours: surface vertical vorticity

Trajectories analyzed
until they reach $z_0 = 50 \text{ m}$

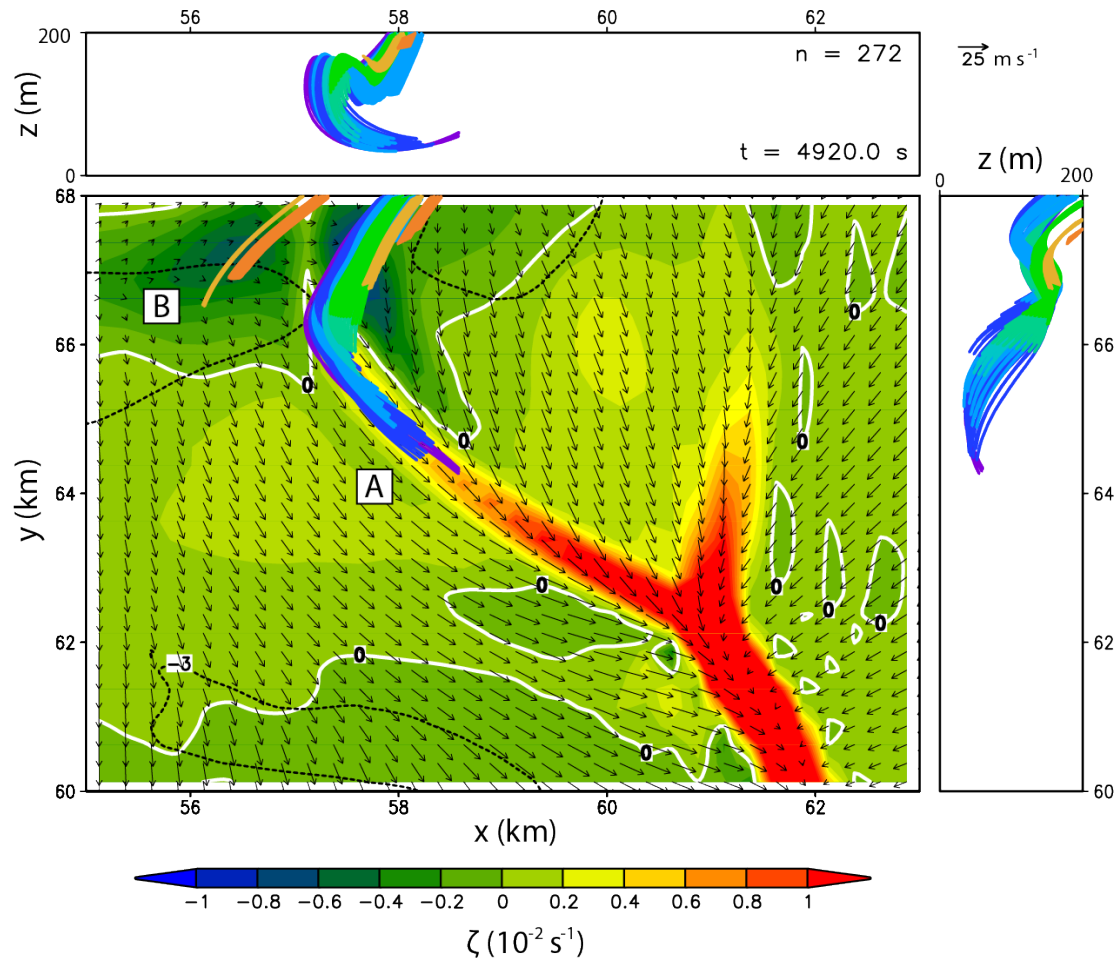
“Rivers” of vertical vorticity in outflow



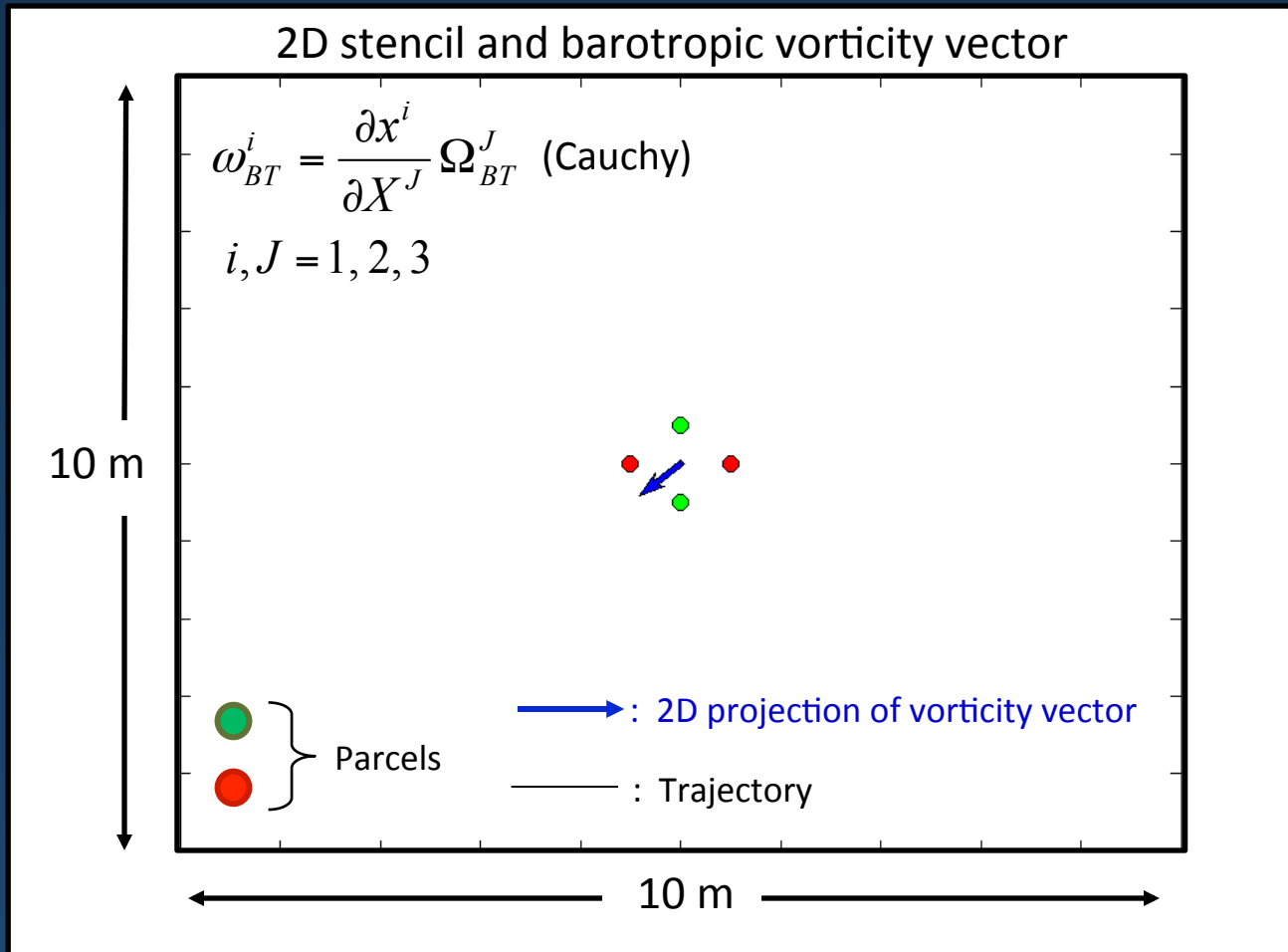
“Rivers” of vertical vorticity in outflow



Forward trajectories and vorticity rivers



Example of barotropic vorticity frozen into a stencil

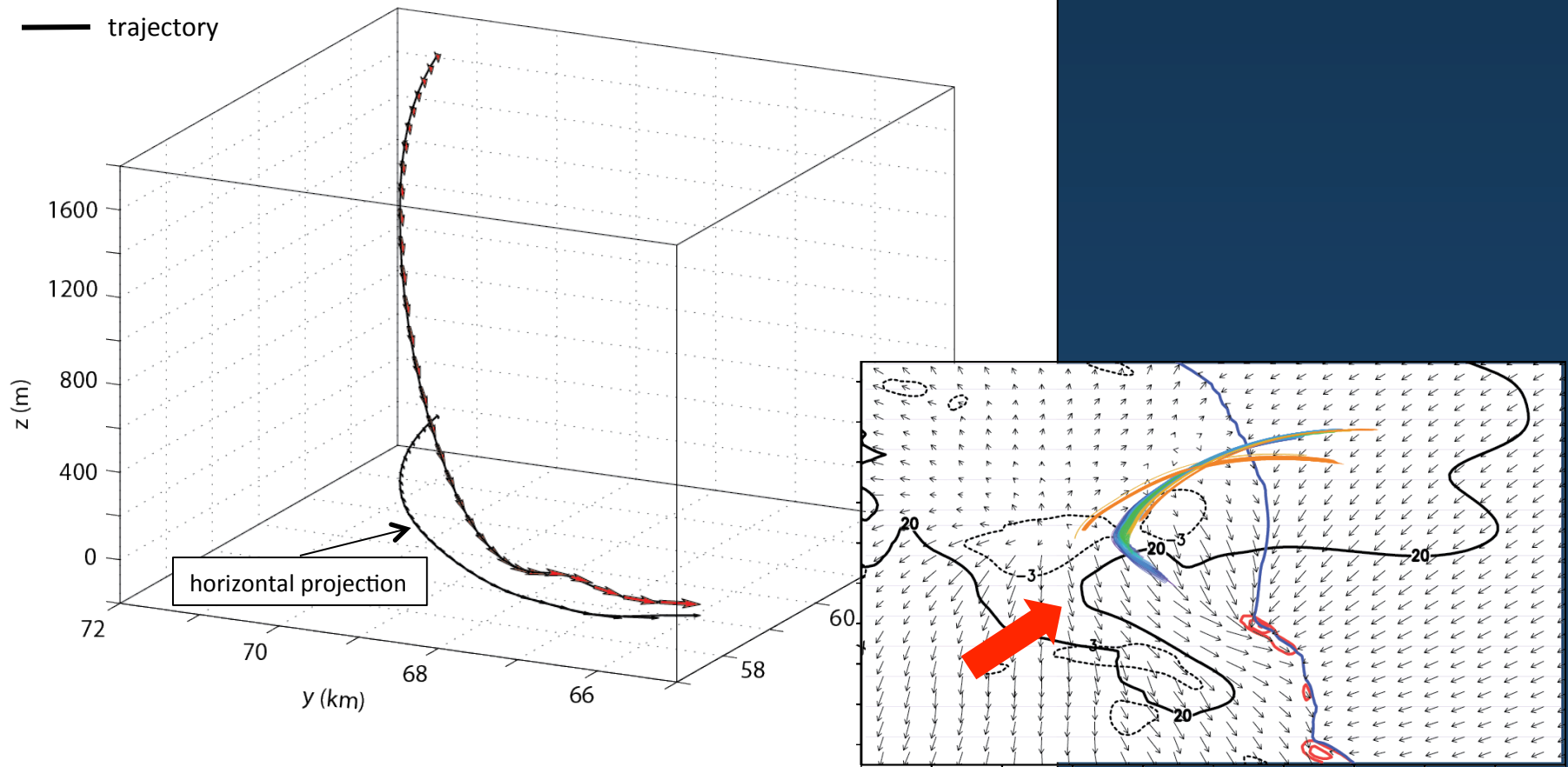


Barotropic vorticity

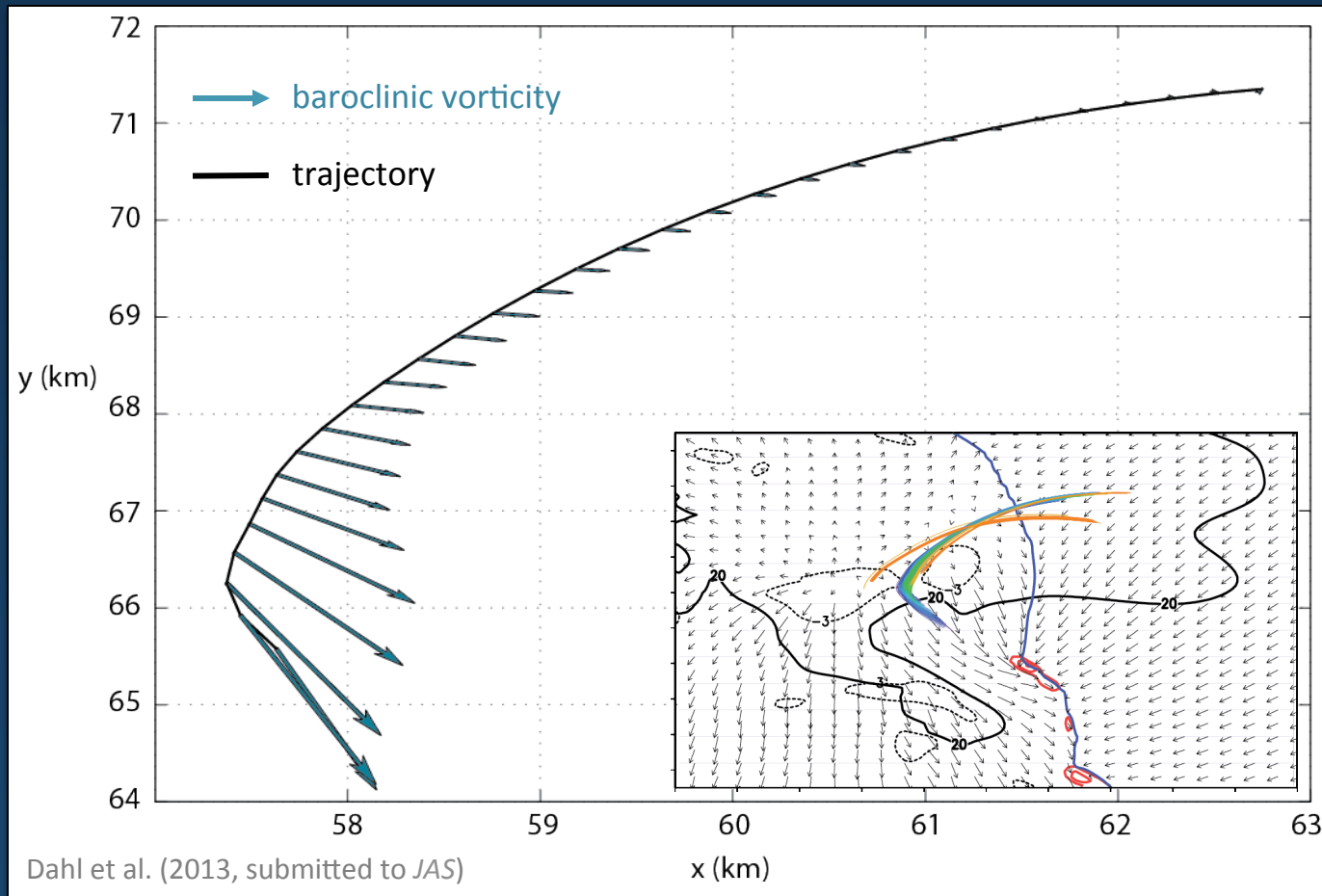
→ barotropic vorticity

— trajectory

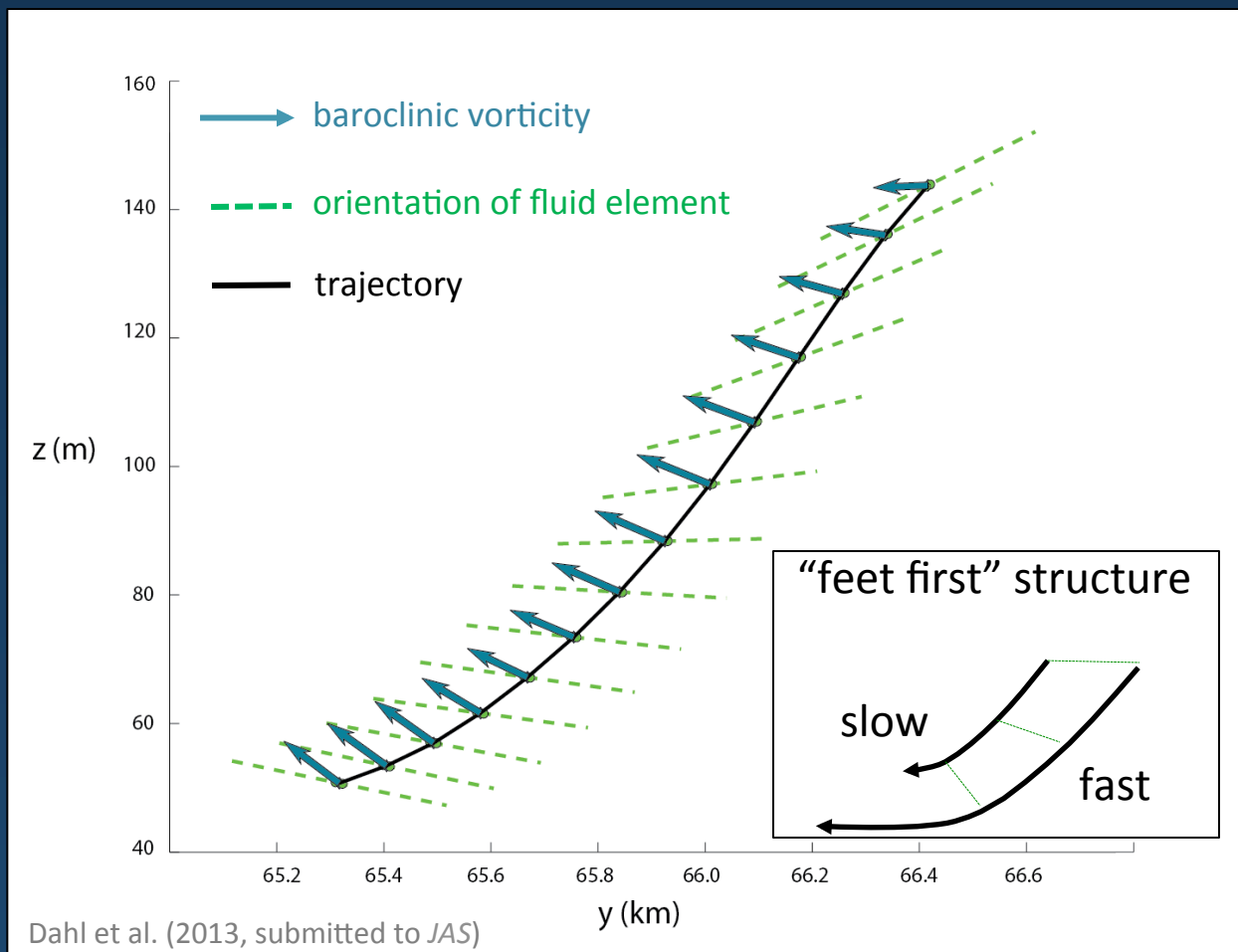
Dahl et al. (2013, submitted to JAS)



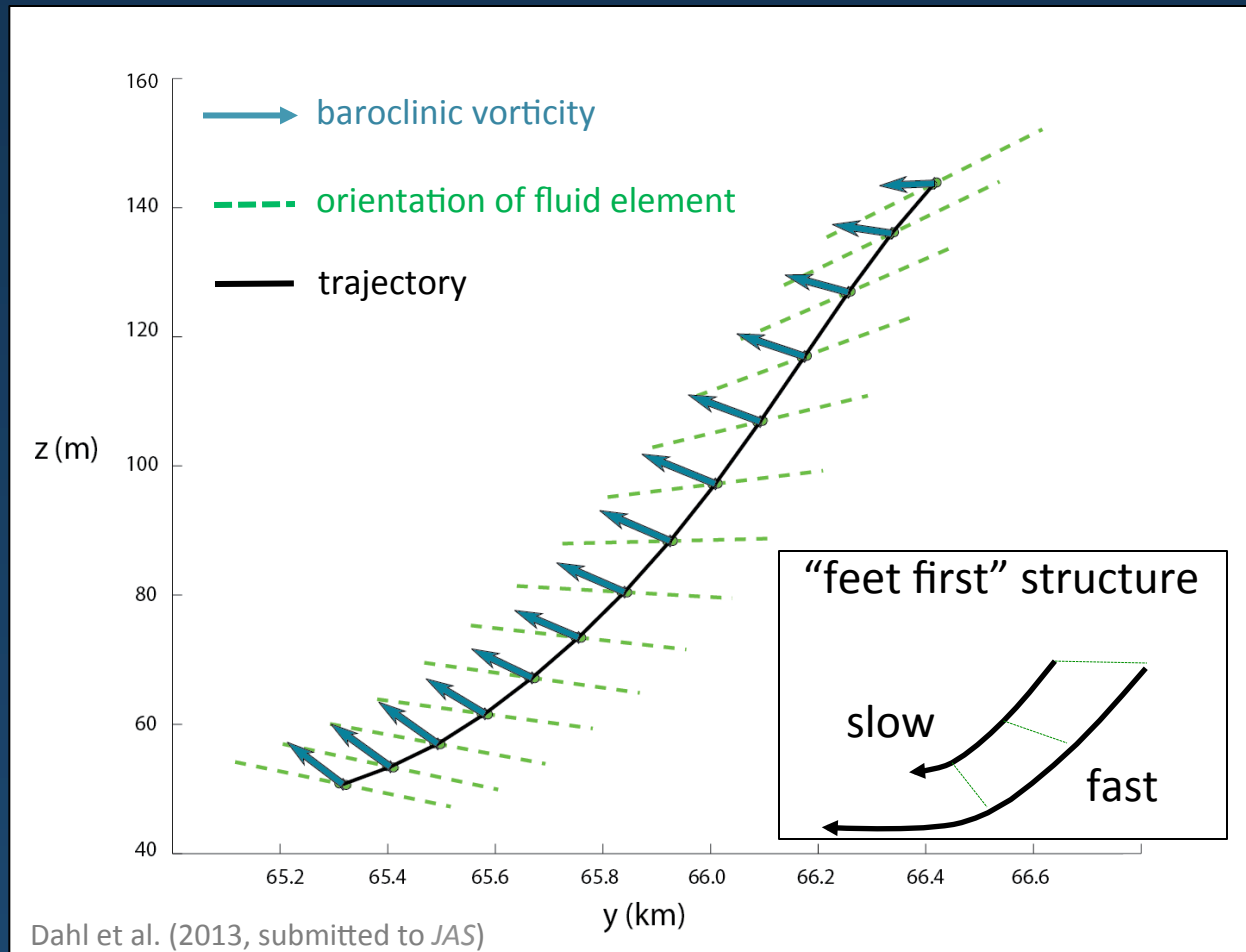
Storm-generated (baroclinic) vorticity



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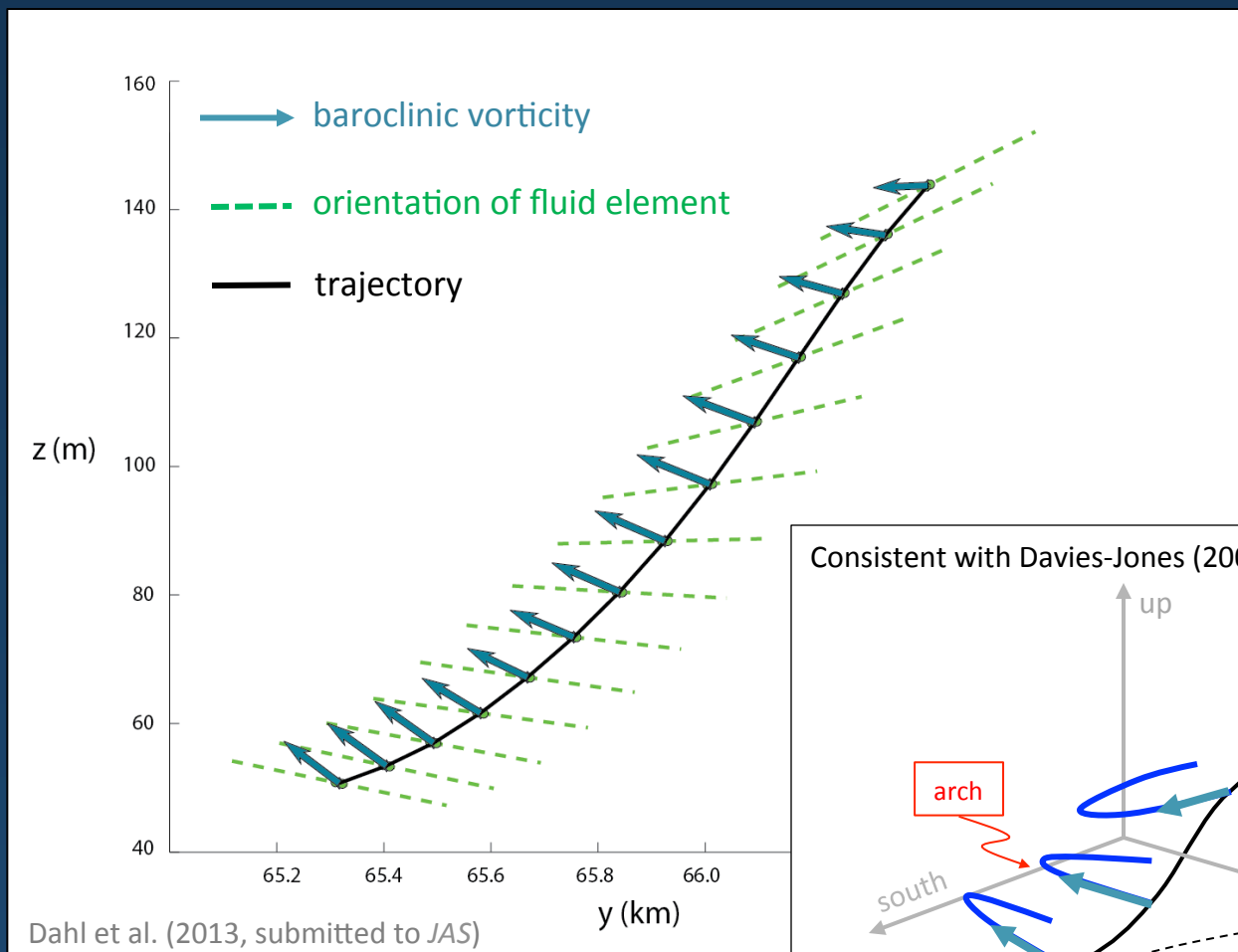


Storm-generated (baroclinic) vorticity

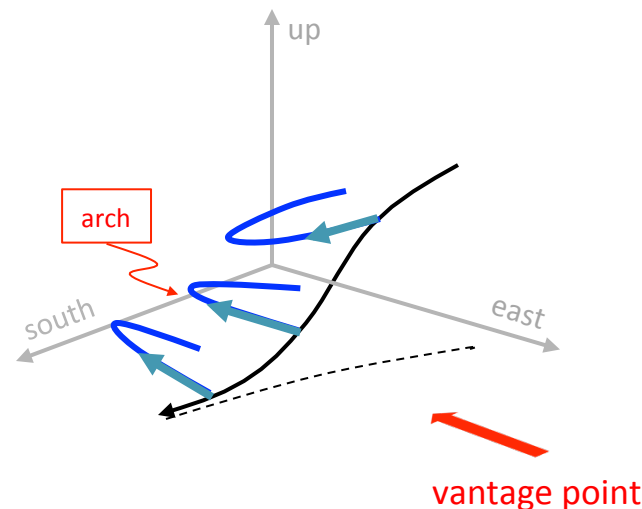


This process was suggested by Davies-Jones and Brooks (1993)

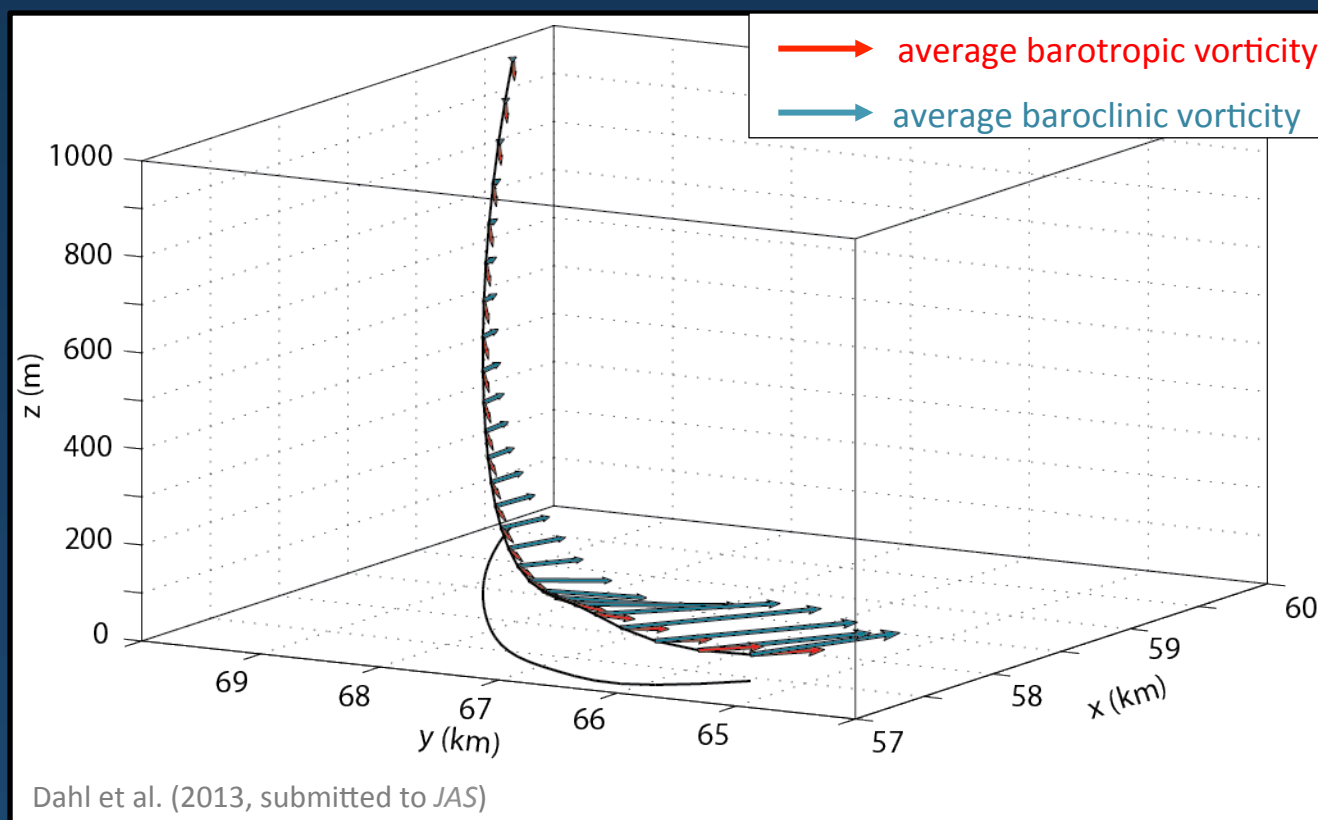
Storm-generated (baroclinic) vorticity



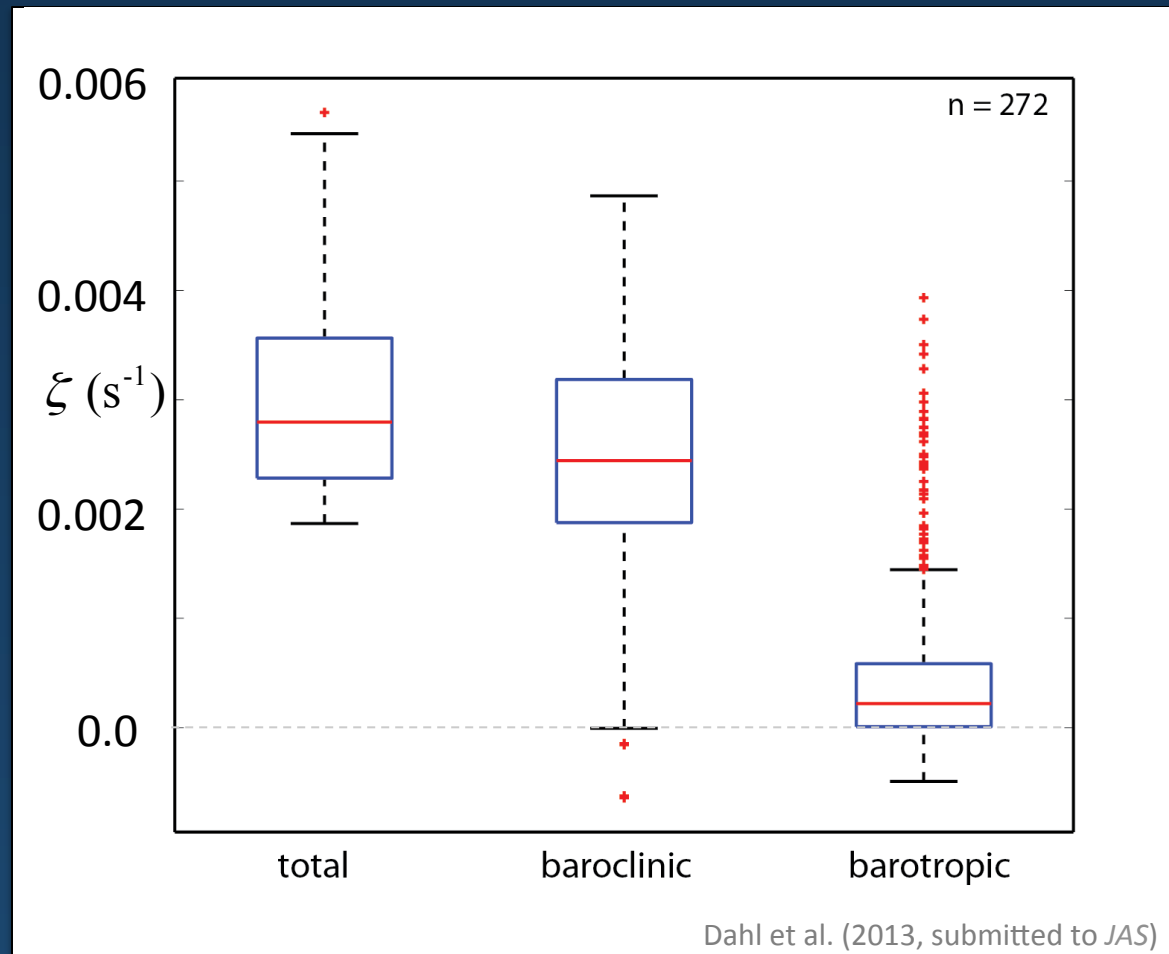
Consistent with Davies-Jones (2000), Straka et al. (2007)



Average over all trajectories: barotropic and baroclinic vorticity



Vertical-vorticity distributions

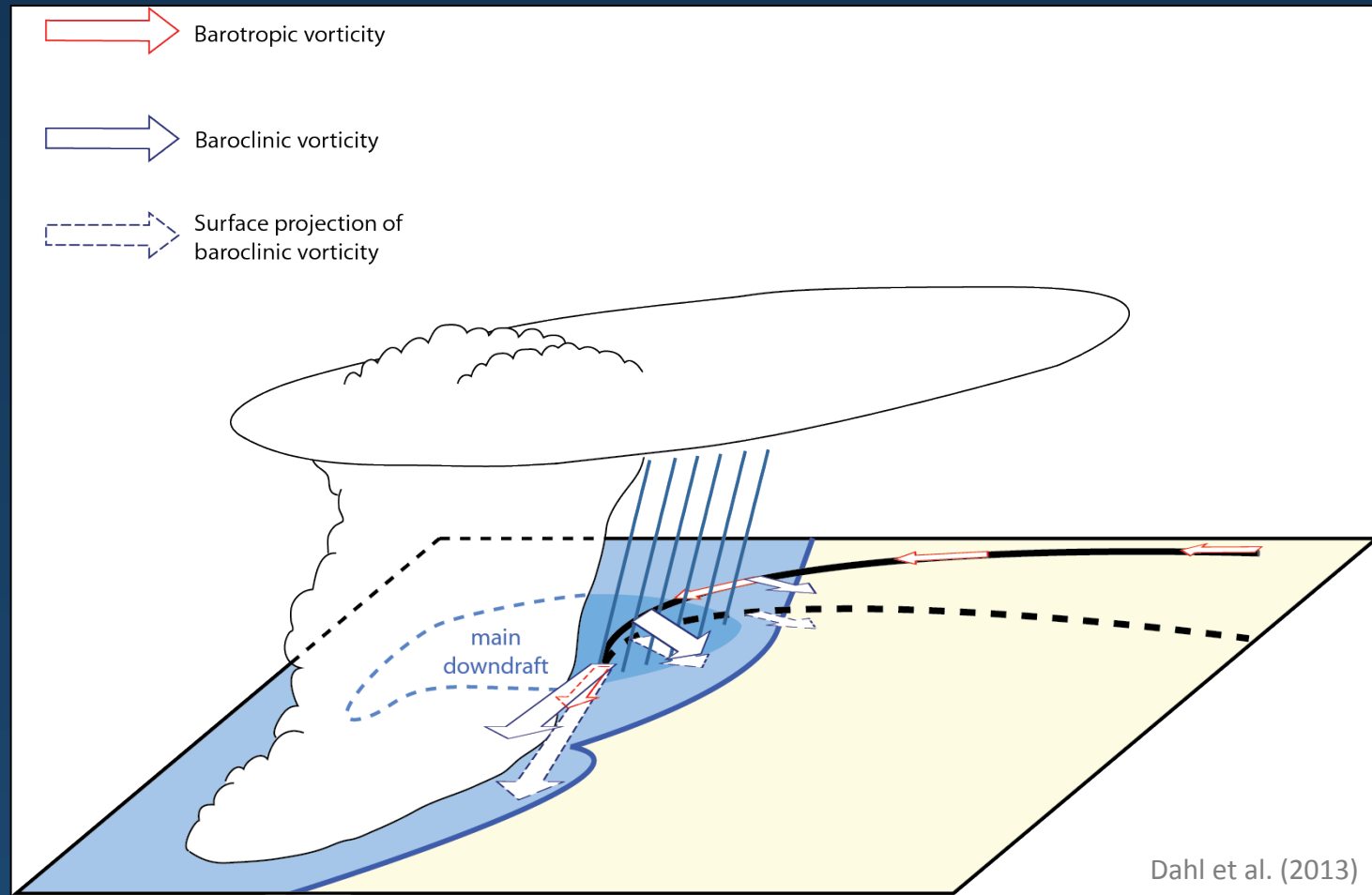


Conclusions

- In this simulation, the barotropic vorticity does not contribute as much to near-ground as the baroclinic vorticity.
- It is likely that the importance of the barotropic vorticity varies based on the orientation of the ambient vorticity:
 - wind shear important at altitudes where downdraft air originates?
- The development of ground-level rotation in supercells is an outflow phenomenon (if there is no vertical vorticity in the environment).

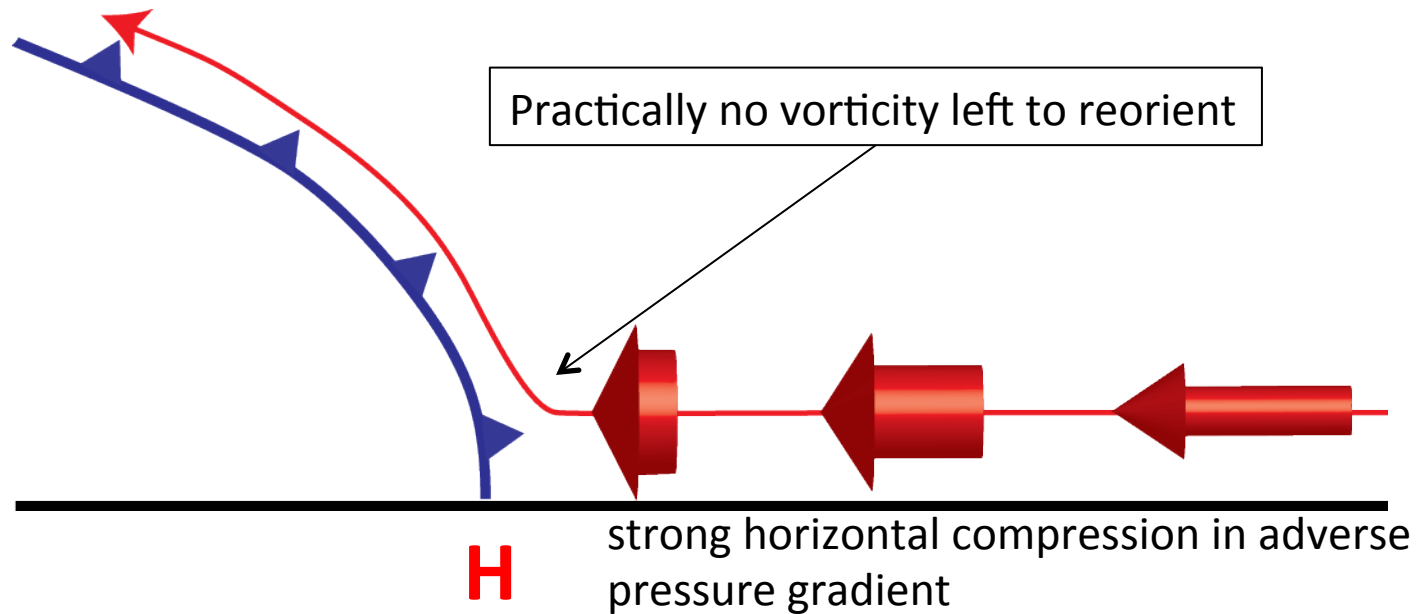
Supplemental slides

Summary

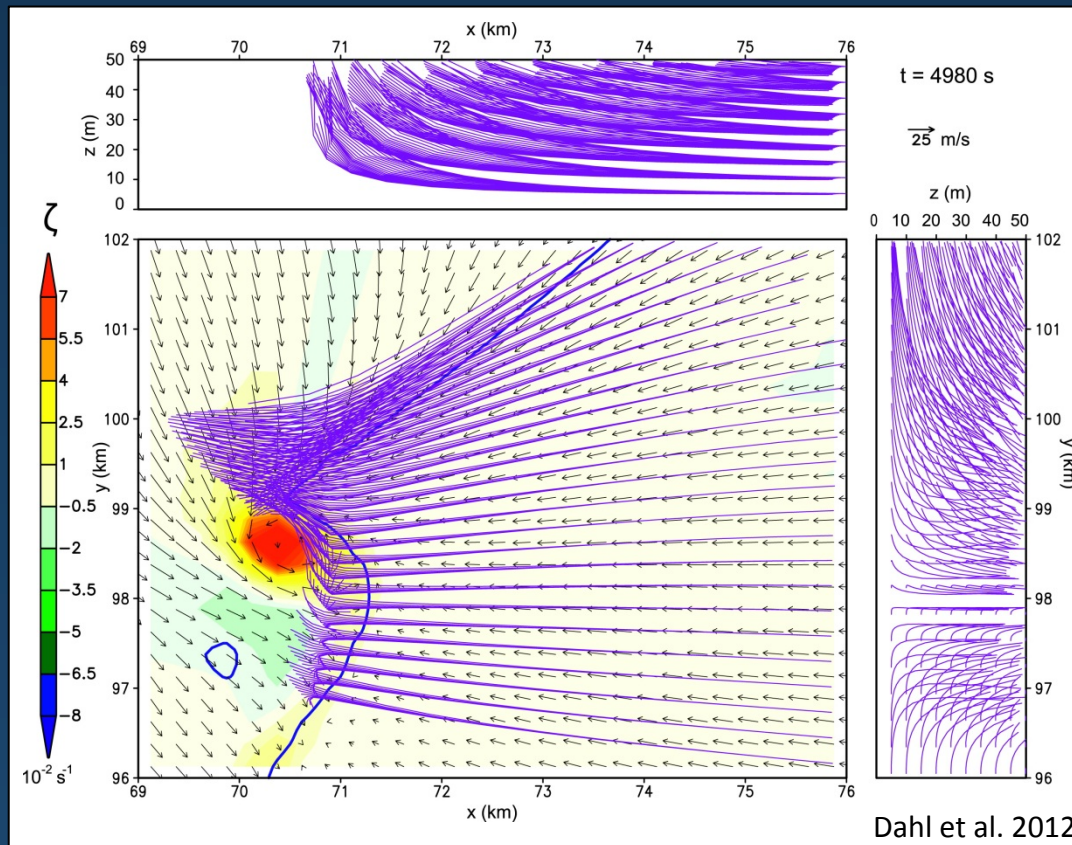


Can a strong gustfront do the trick?

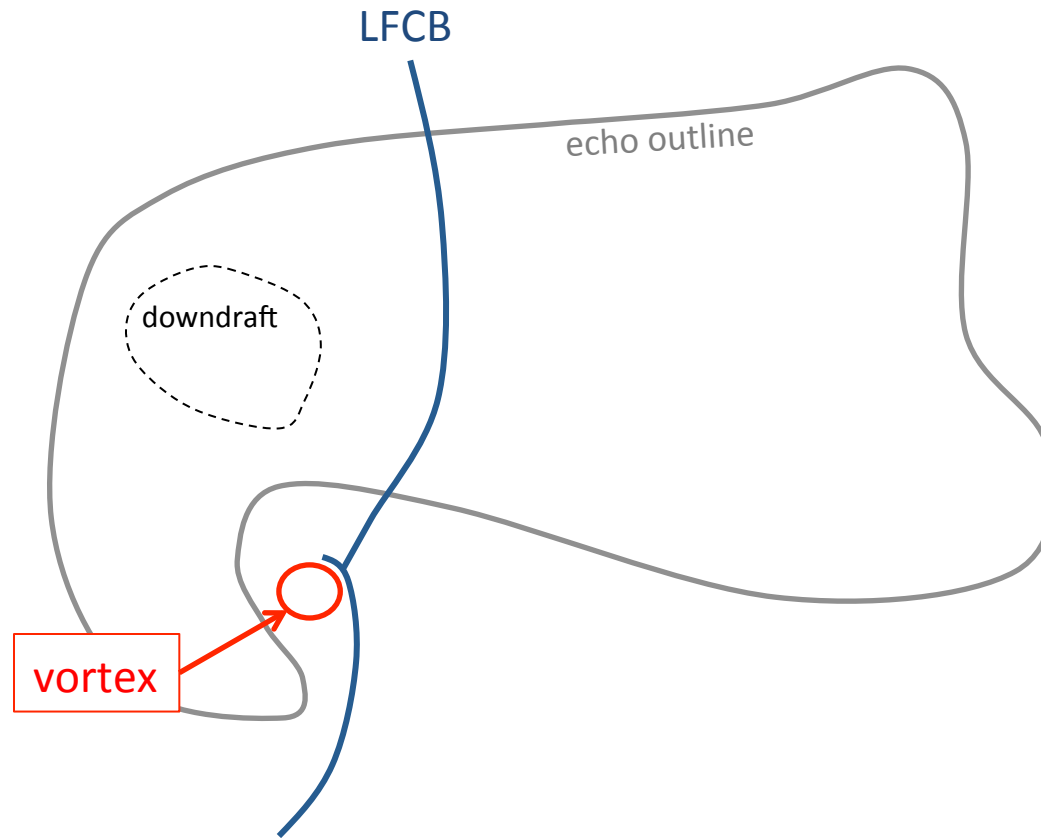
After Davies-Jones and Markowski 2013



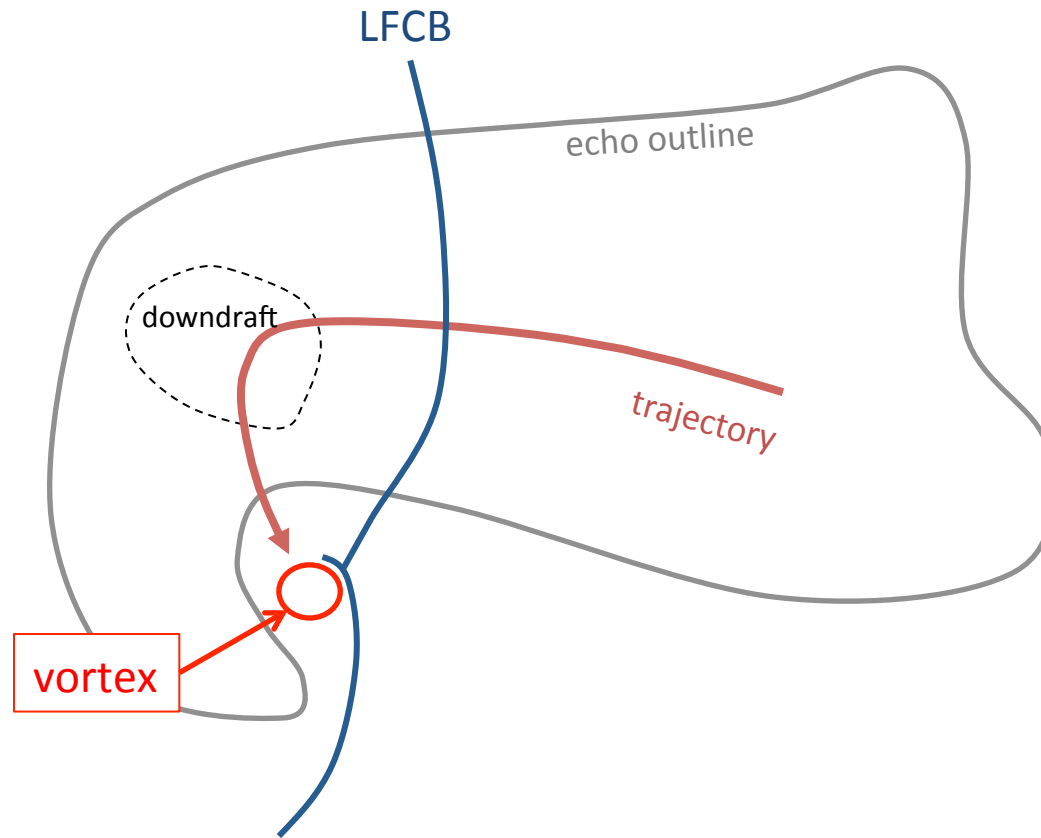
“Inflow” parcels



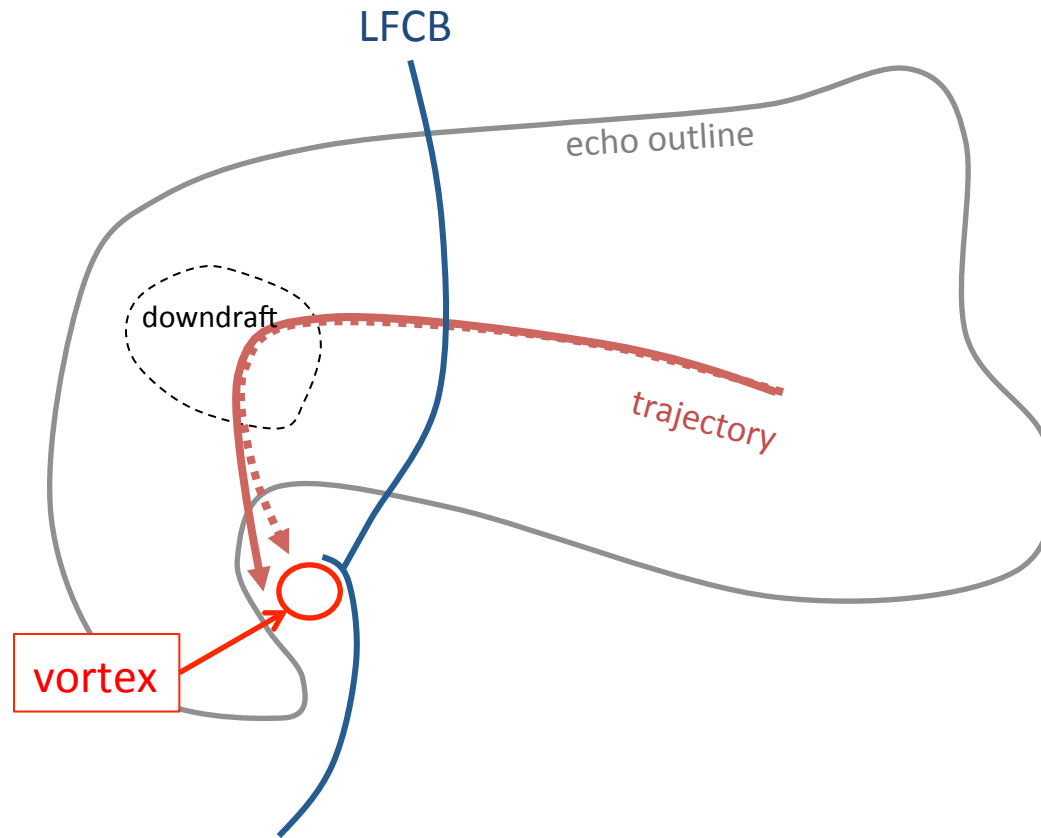
Forward trajectory errors?




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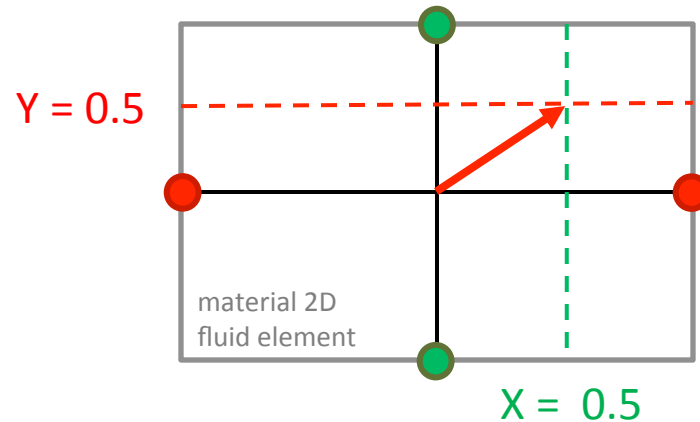


Barotropic vorticity: Frozen into the medium


frozen vorticity vector

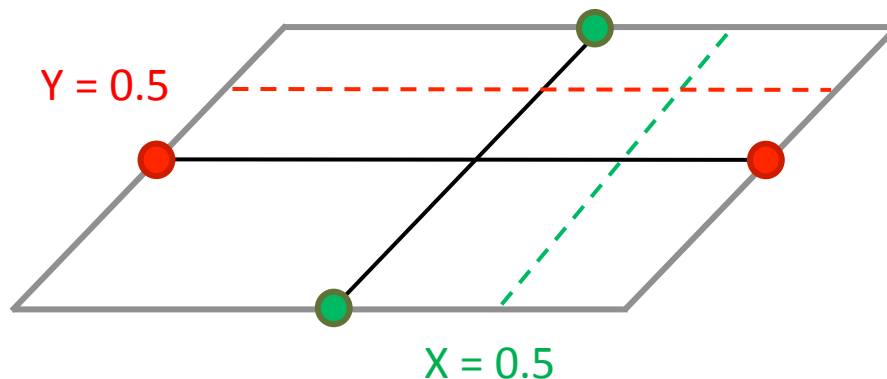
  } parcels of air

Reference configuration at t_0
(Lagrangian = Eulerian frame)



Barotropic vorticity: Frozen into the medium

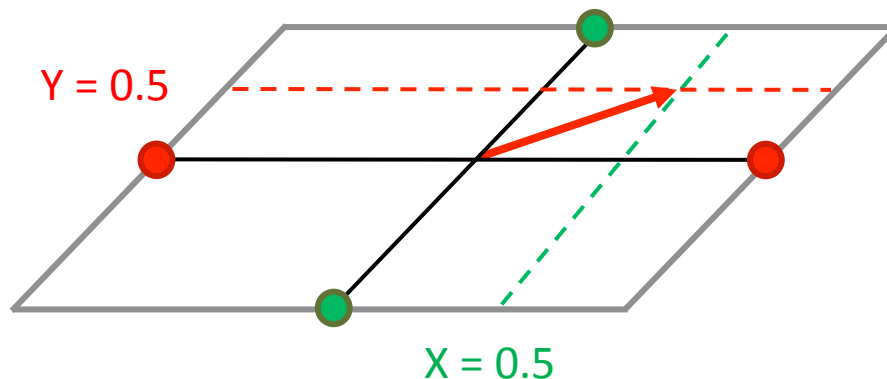
Configuration at t
(Lagrangian frame)



Barotropic vorticity: Frozen into the medium


frozen vorticity vector

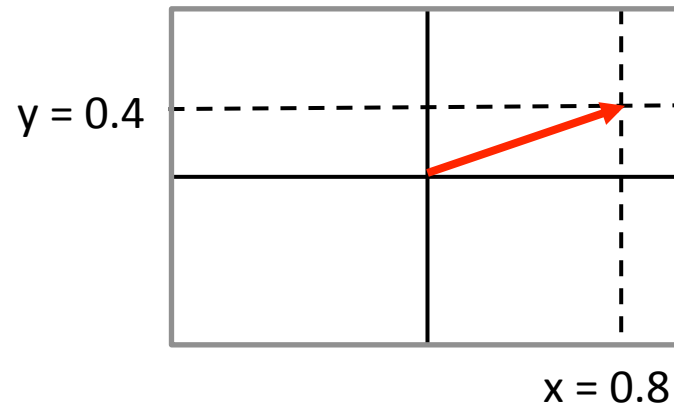
Configuration at t
(Lagrangian frame)



Transformation to Cartesian grid


frozen vorticity vector

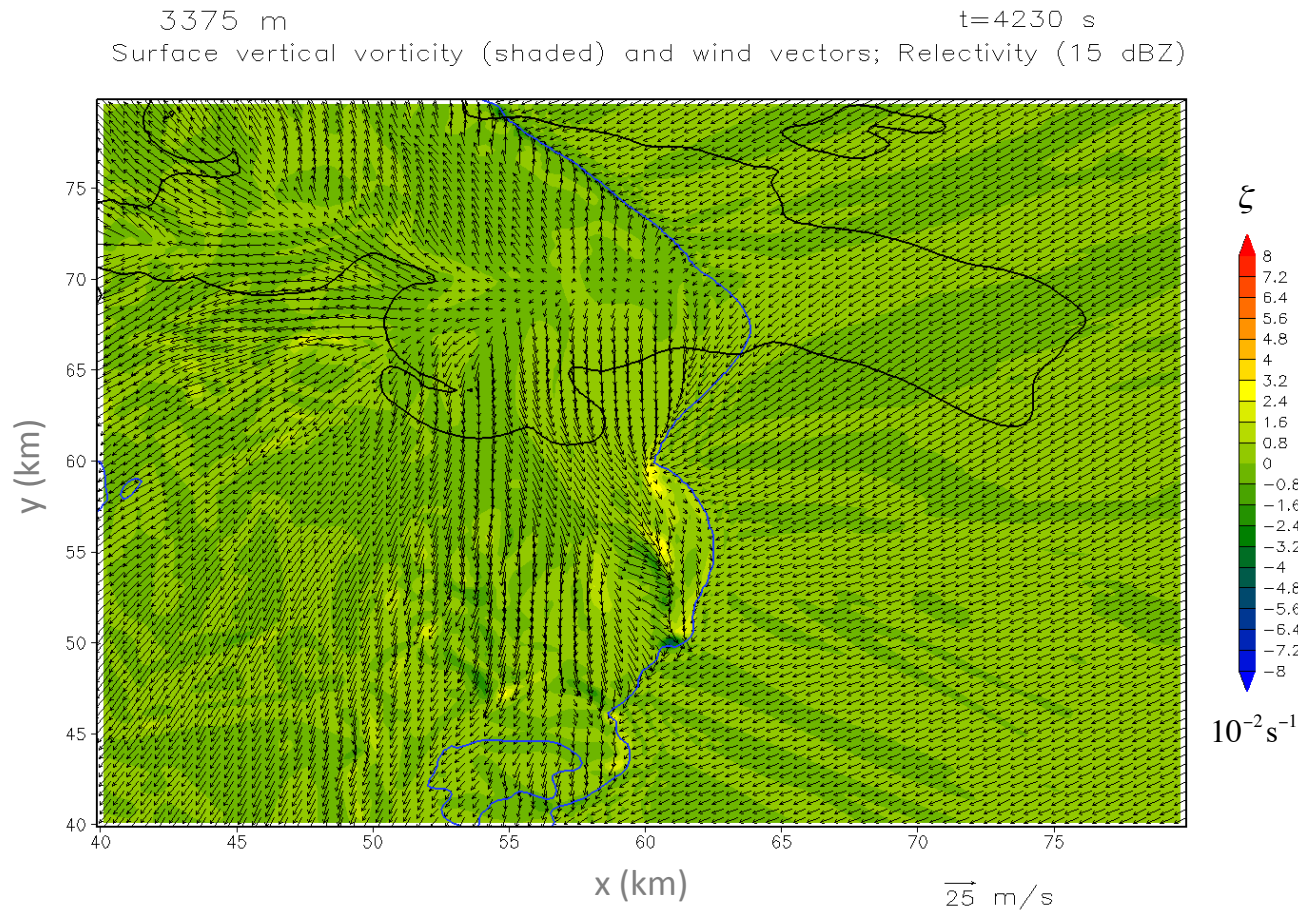
Configuration at t
(Eulerian frame)



A coordinate transformation yields the barotropic vorticity in spatial coordinates (Cauchy's formula):

$$\omega_{BT}^i = \frac{\partial x^i}{\partial X^J} \Omega_{BT}^J \quad i, J = 1, 2, 3$$

Surface vertical vorticity



Storm-generated (baroclinic) vorticity

