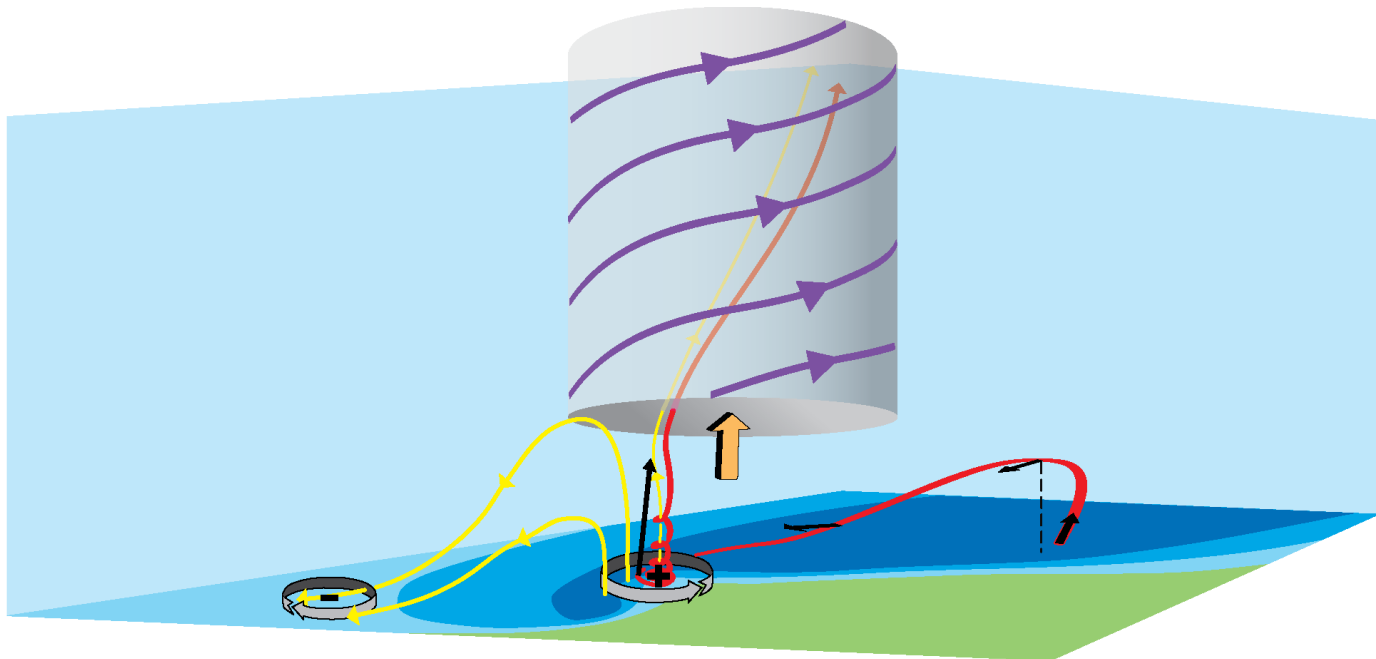


The influence of environmental low-level shear and cold pools on tornadogenesis: Insights from toy simulations

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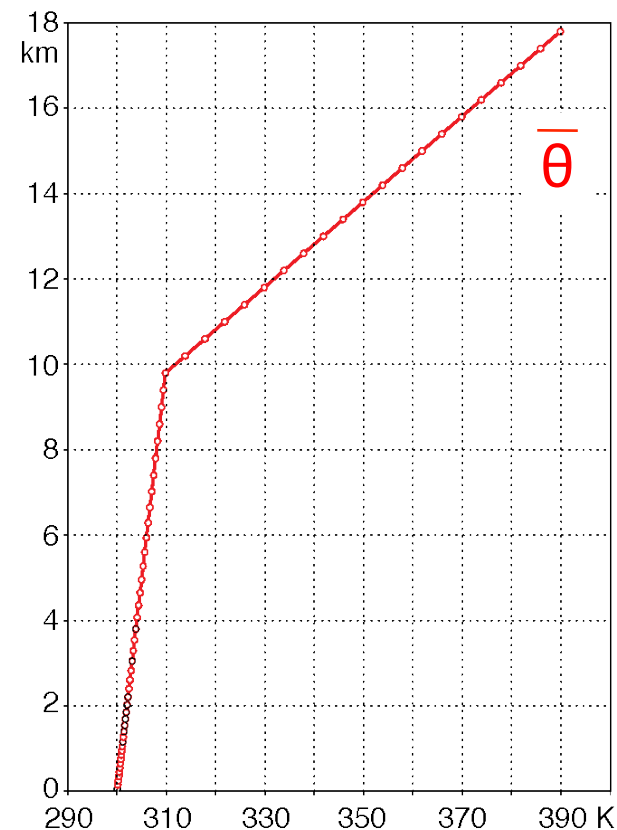


Motivating Questions

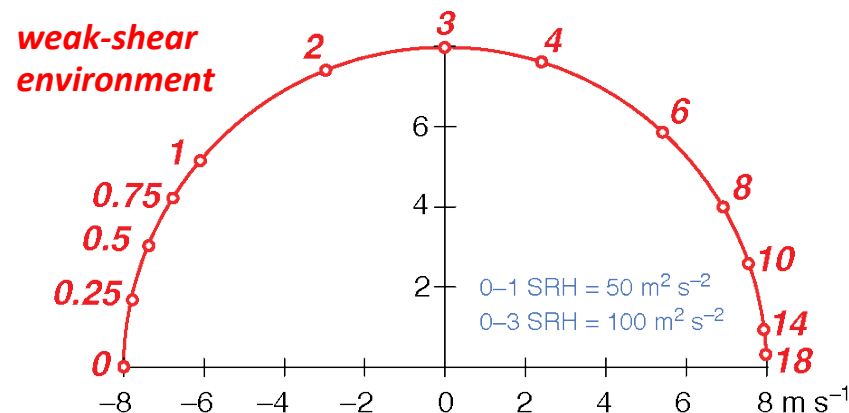
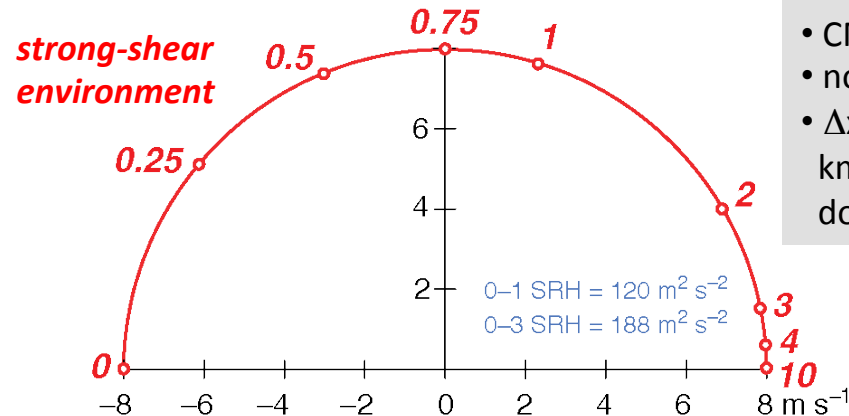
- What is the role of environmental horizontal vorticity in tornadogenesis?
 - Large environmental horizontal vorticity at low levels favors tornadic supercells
 - However, there's a mountain of evidence from both observations and simulations that the *baroclinic mechanism* is the dominant way by which near-surface cyclonic vorticity develops
 - e.g., material circuit analyses, vortex line analyses, and barotropic vs baroclinic vorticity decompositions
- Why is the combination of strong vertical shear and high boundary-layer RH so favorable for tornadoes?

The Approach: “A Toy Model”

- Drive an updraft with a stationary, cylindrical heat source
 - The environmental wind profile is a semicircle (purely streamwise environmental vorticity); the updraft rotates cyclonically, rotation is a maximum at midlevels, rotation vanishes at the surface.



Walko (1993)

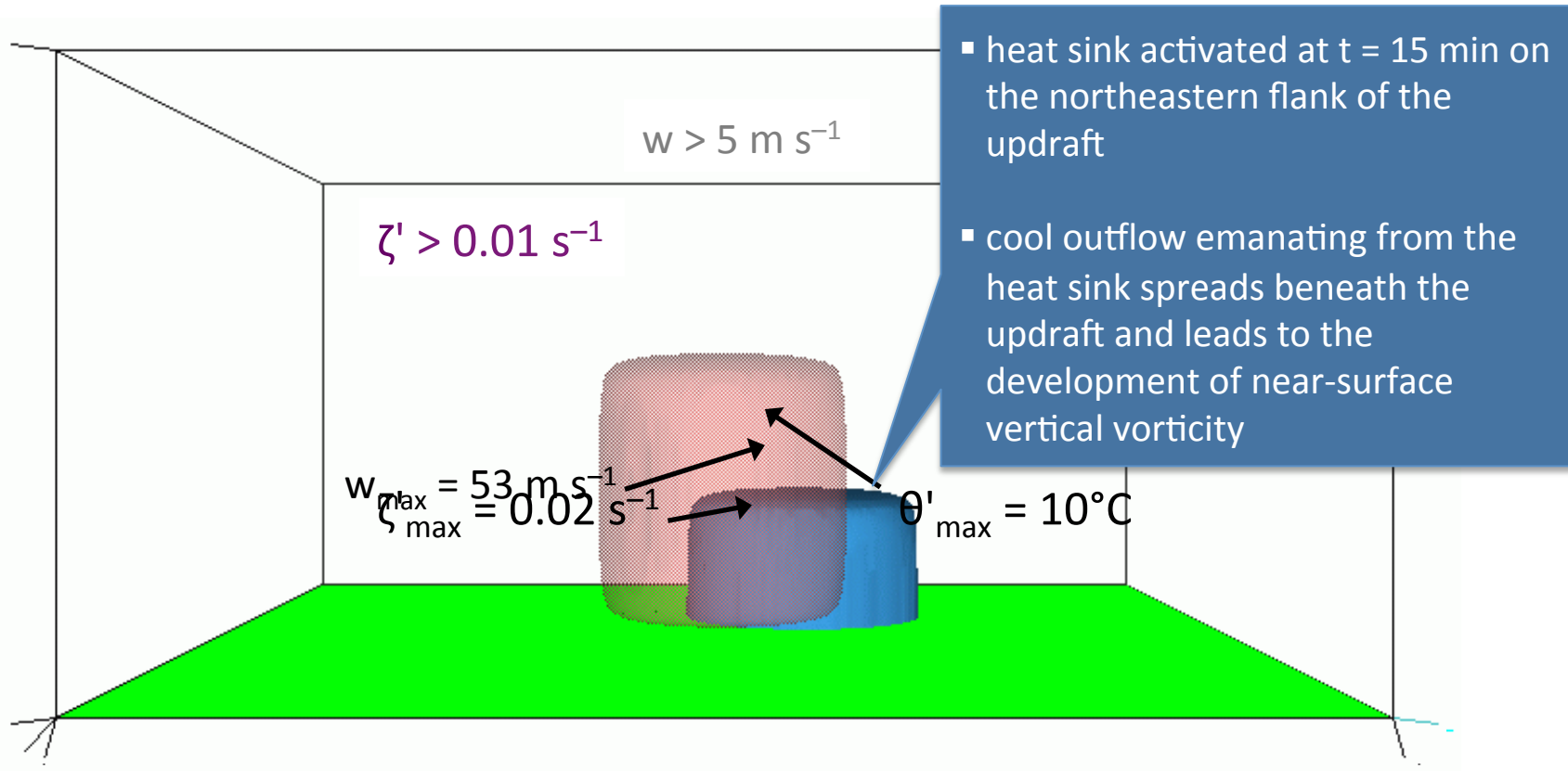


McCaul & Weisman (2001)

- CM1
- no water vapor
- $\Delta x, \Delta y = 100 \text{ m}$ in $25 \times 25 \text{ km}$ inner domain (entire domain is $100 \times 100 \text{ km}$)

The Approach: “A Toy Model”

The simulations are not as realistic as those that incorporate more physical processes (e.g., latent heating, microphysics, etc.), but we believe the simplicity of such “toy models” can be advantageous in that it may be easier to isolate key dynamical processes and assign cause and effect.

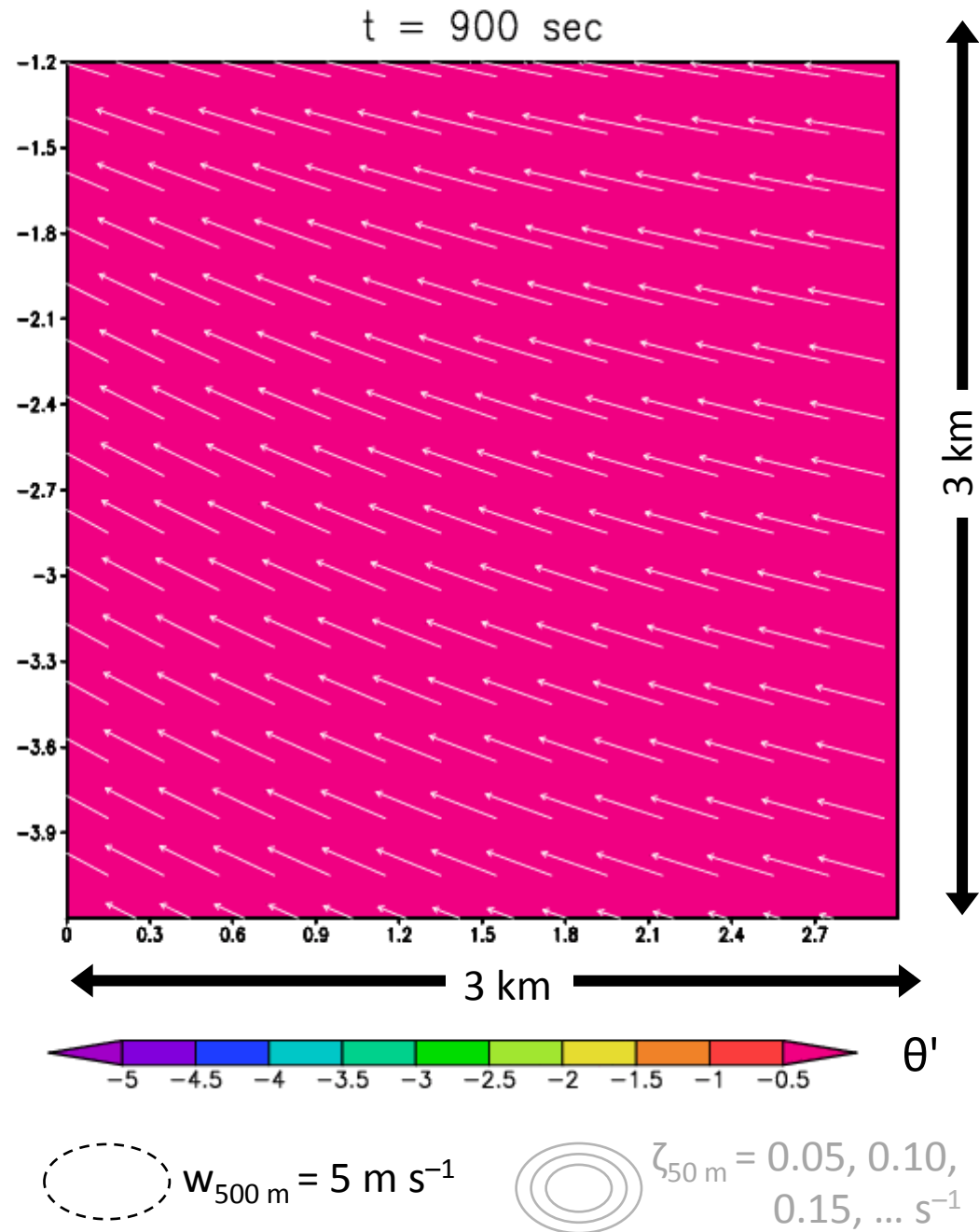
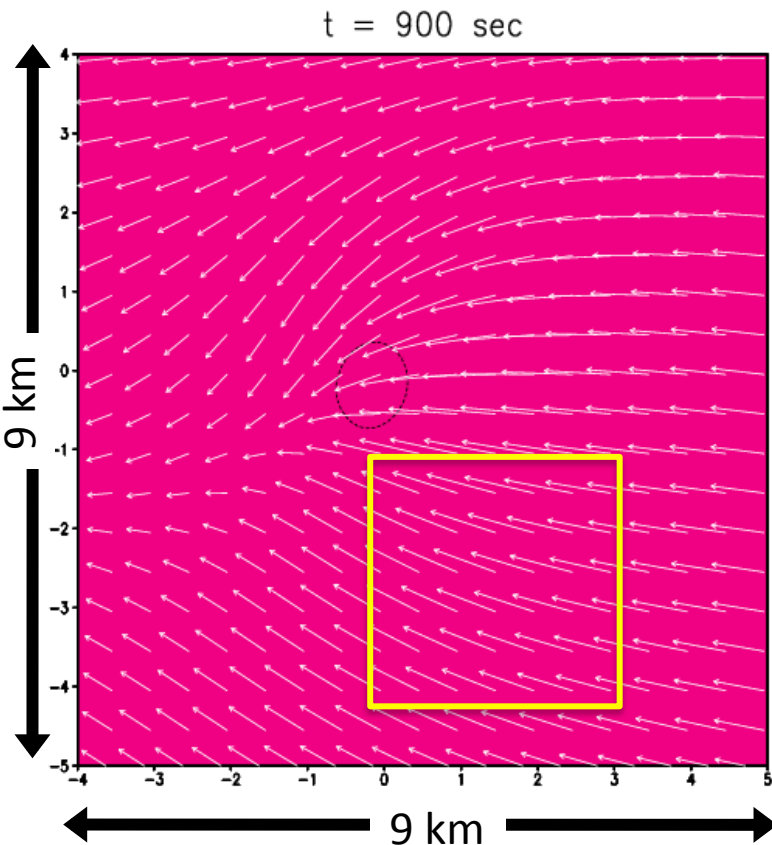


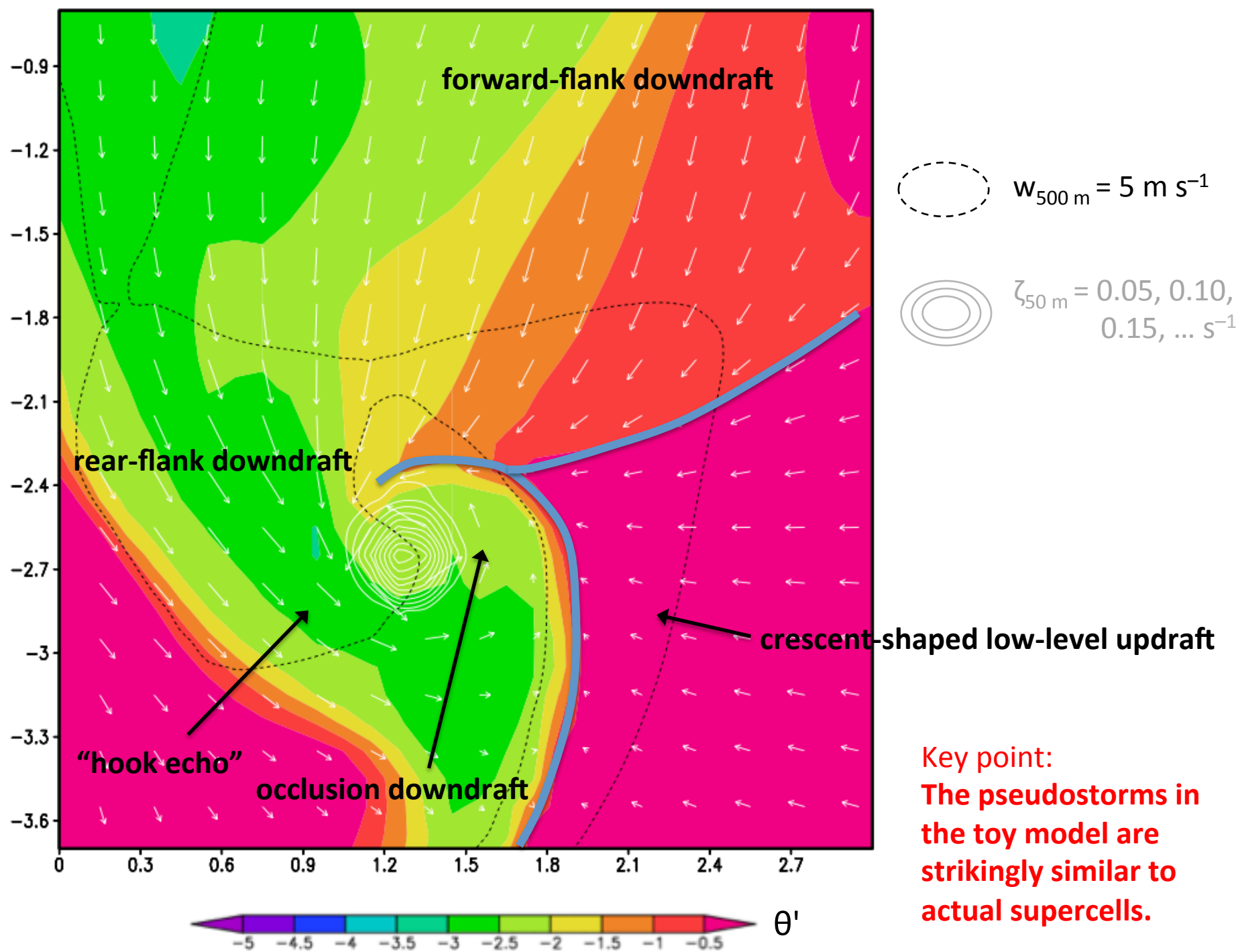
**strong low-level shear,
intermediate heat sink**

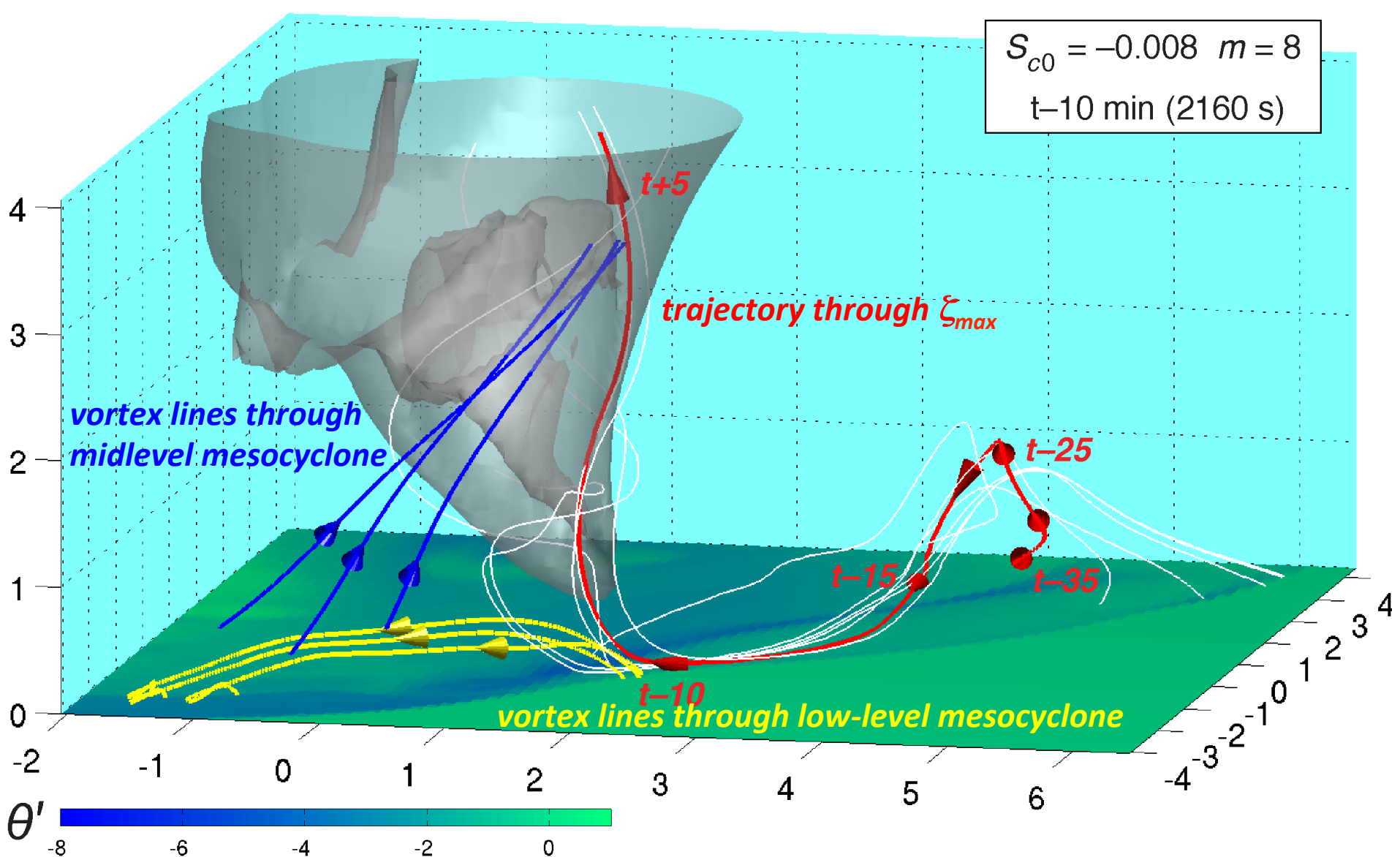
$$\zeta_{\max} = 0.83 \text{ s}^{-1}$$

$$v_{\max} = 31 \text{ m s}^{-1}$$

$$\text{duration} = 9.4 \text{ min}$$



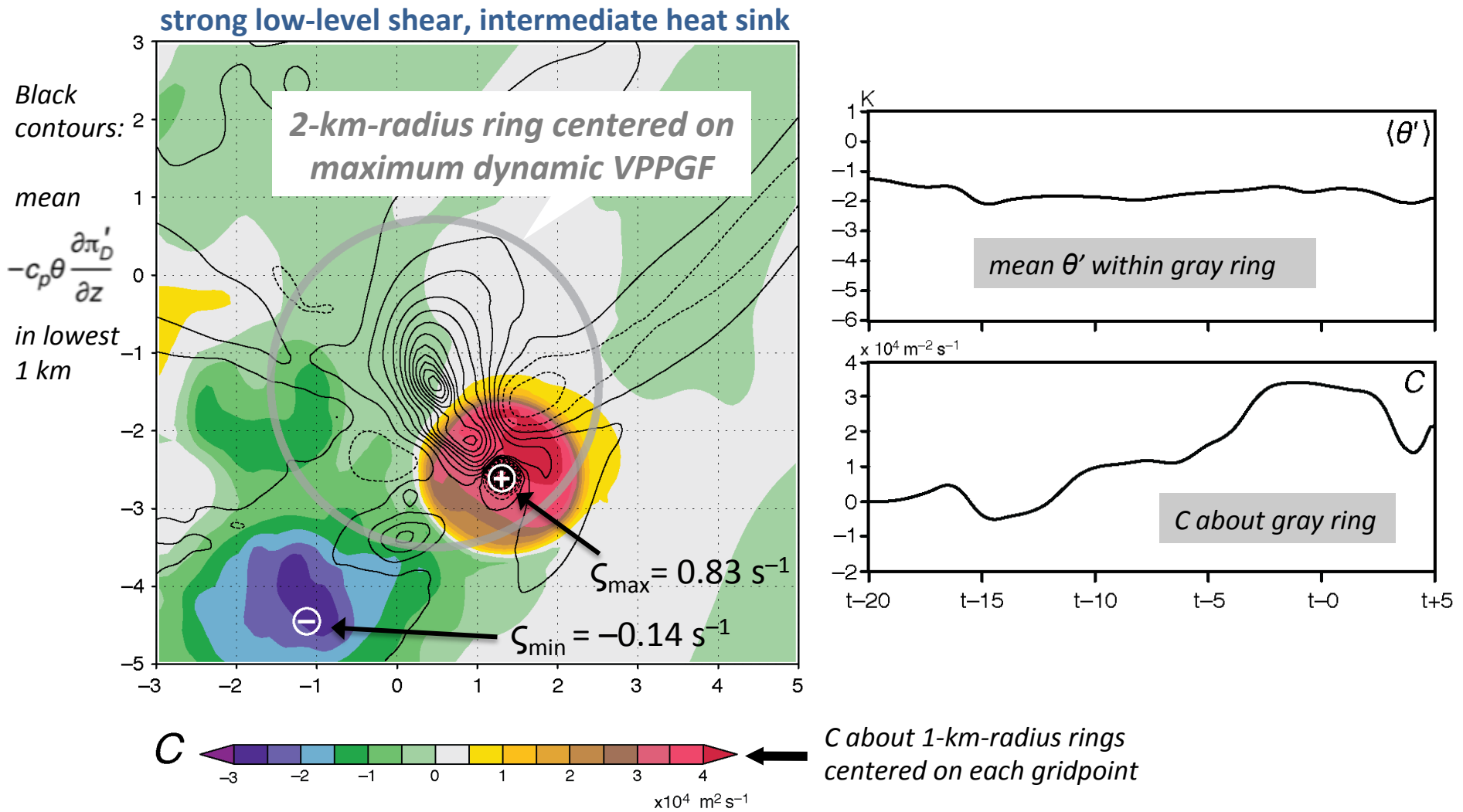




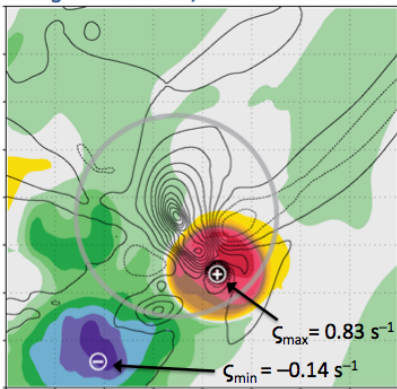
Key point:

There's "higher-order" similarity of the pseudostorms to actual supercells as well (e.g., vortex lines, trajectories, and the dynamical processes leading to the development of near-surface rotation compare well with actual supercells despite the simplicity of the toy model).

The development of an intense cyclonic vortex is the result of **circulation-rich, near-surface air being associated with weak negative buoyancy** and also experiencing **a large upward-directed VPPGF** owing to its proximity to the midlevel mesocyclone.

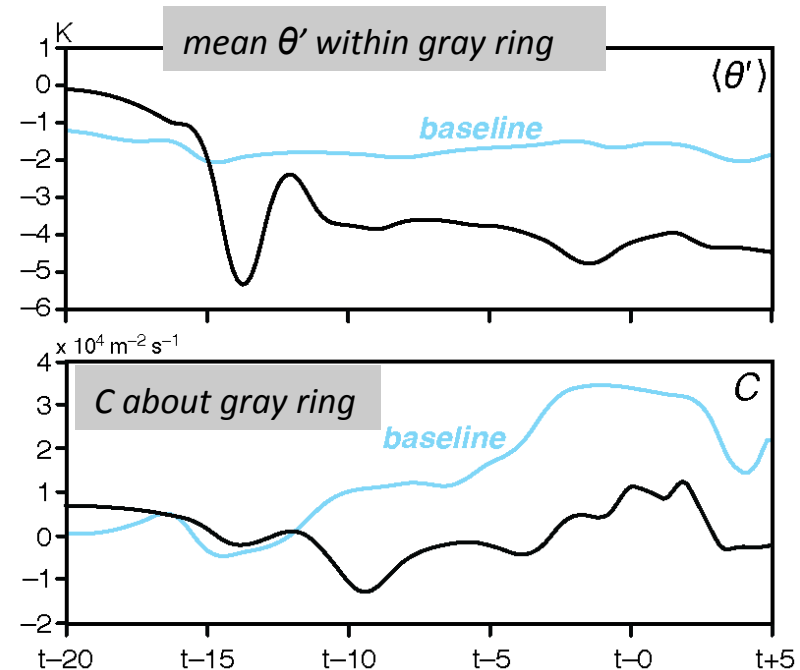
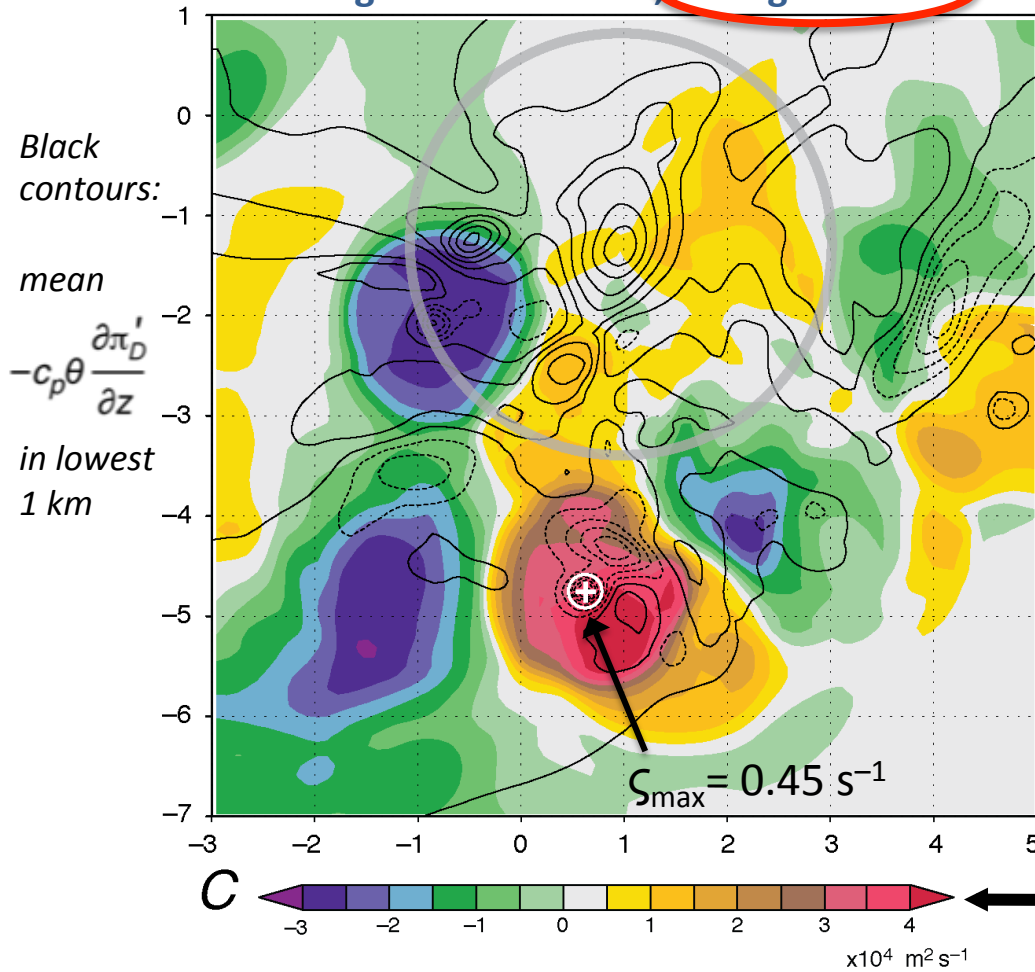


strong low-level shear, intermediate heat sink

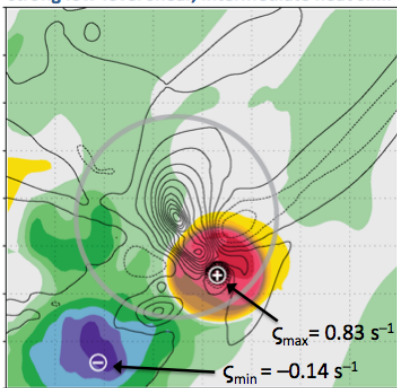


When the heat sink is strong, the maximum near-surface circulation is (slightly) stronger, but **it is displaced farther from the region of dynamic lifting and is associated with larger negative buoyancy.**

strong low-level shear, **strong heat sink**

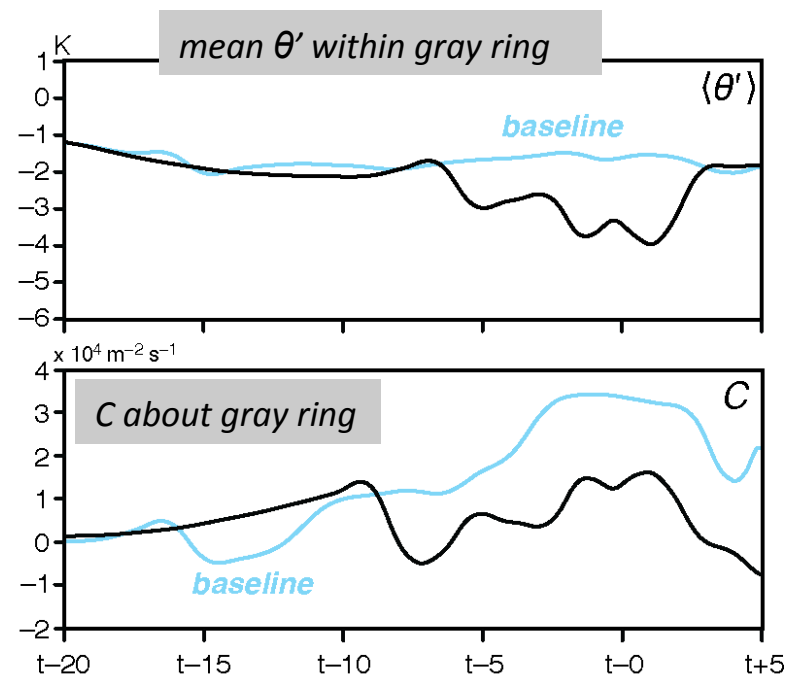
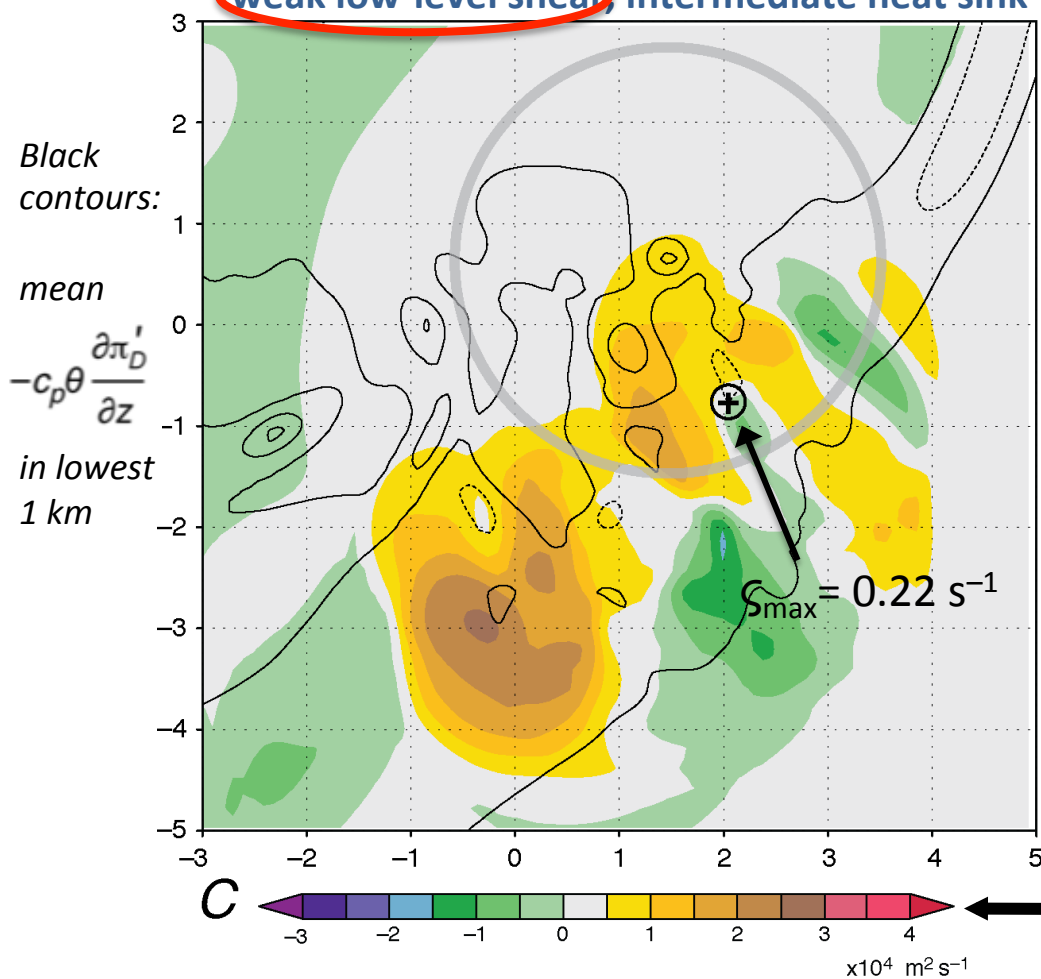


C about 1-km-radius rings
 centered on each gridpoint



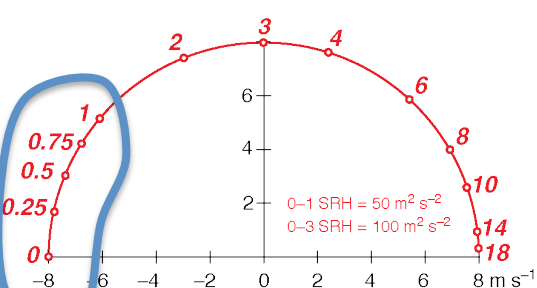
When the low-level shear is weak, **the dynamic lifting is much weaker.** [The near-surface circulation also is less than in the baseline (intense vortex) simulation, displaced far from the dynamic lifting, and is associated with greater negative buoyancy. **This illustrates the complex relationship between parcel residence times in the heat sink, exposure to baroclinity, and the location where large circulation develops relative to the dynamic forcing!**]

weak low-level shear, intermediate heat sink

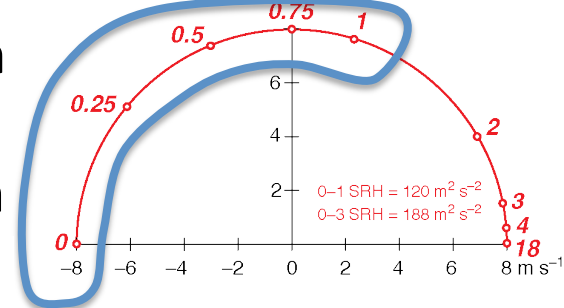


C about 1-km-radius rings centered on each gridpoint

Dynamic VPPGF associated with quasi-steady rotating updraft just prior to heat sink activation

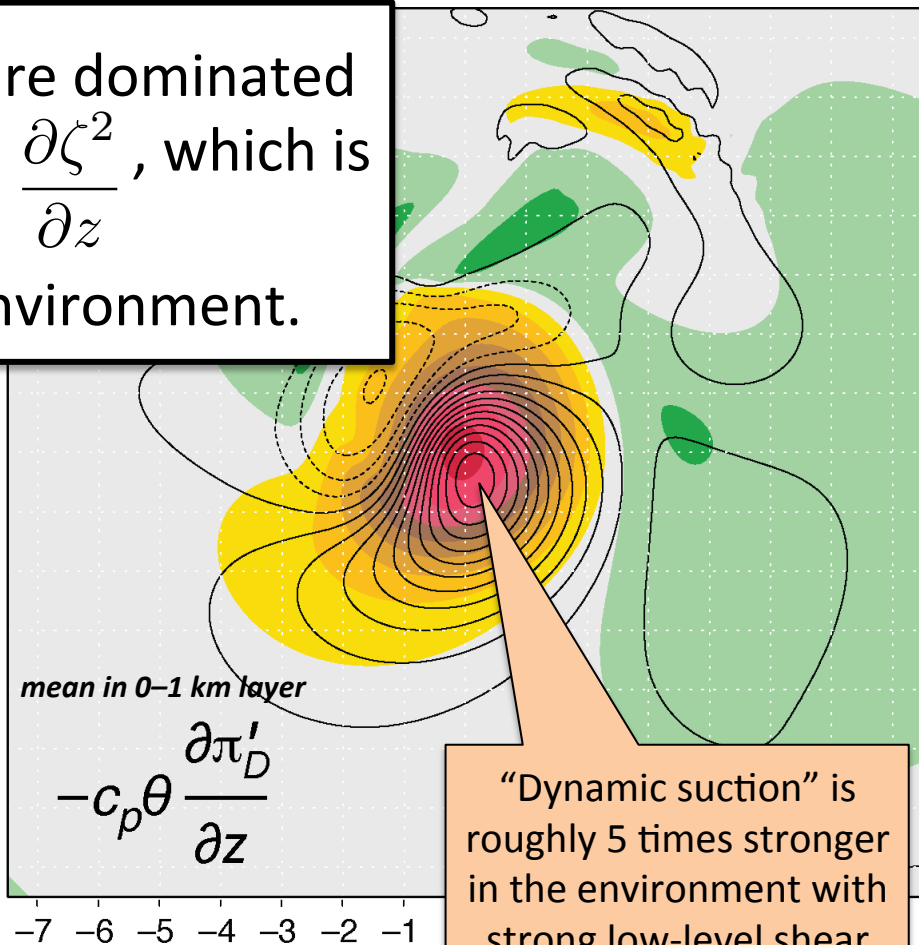
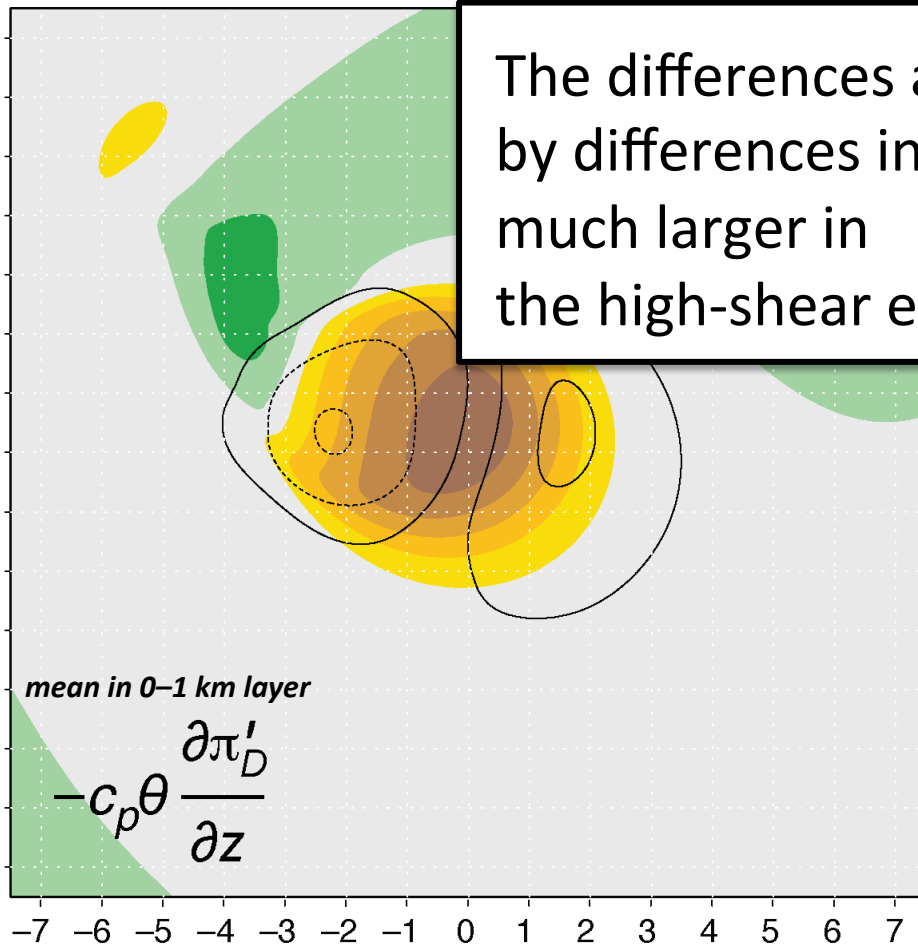


Weak low-level shear

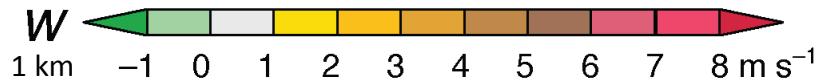


Strong low-level shear

The differences are dominated by differences in $\frac{\partial \zeta^2}{\partial z}$, which is much larger in the high-shear environment.



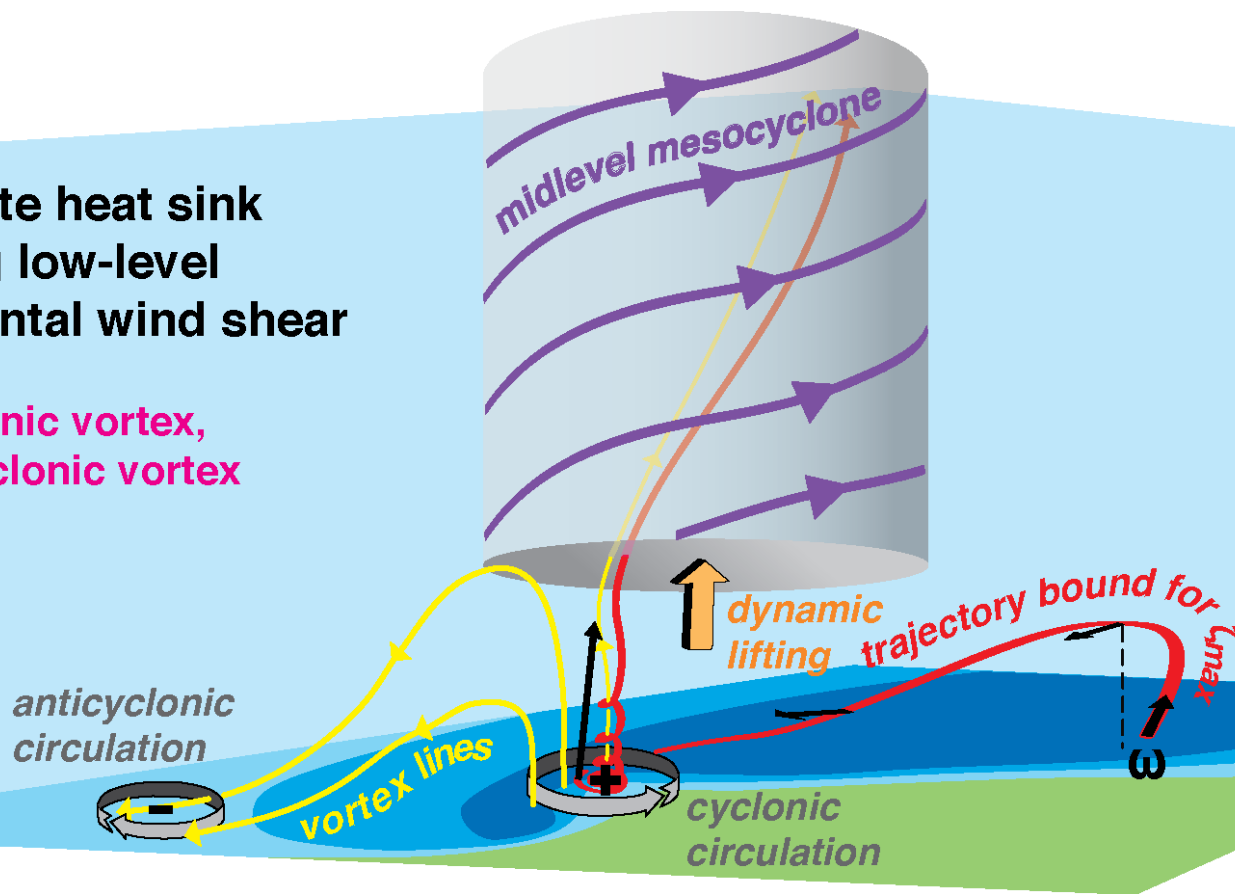
"Dynamic suction" is roughly 5 times stronger in the environment with strong low-level shear



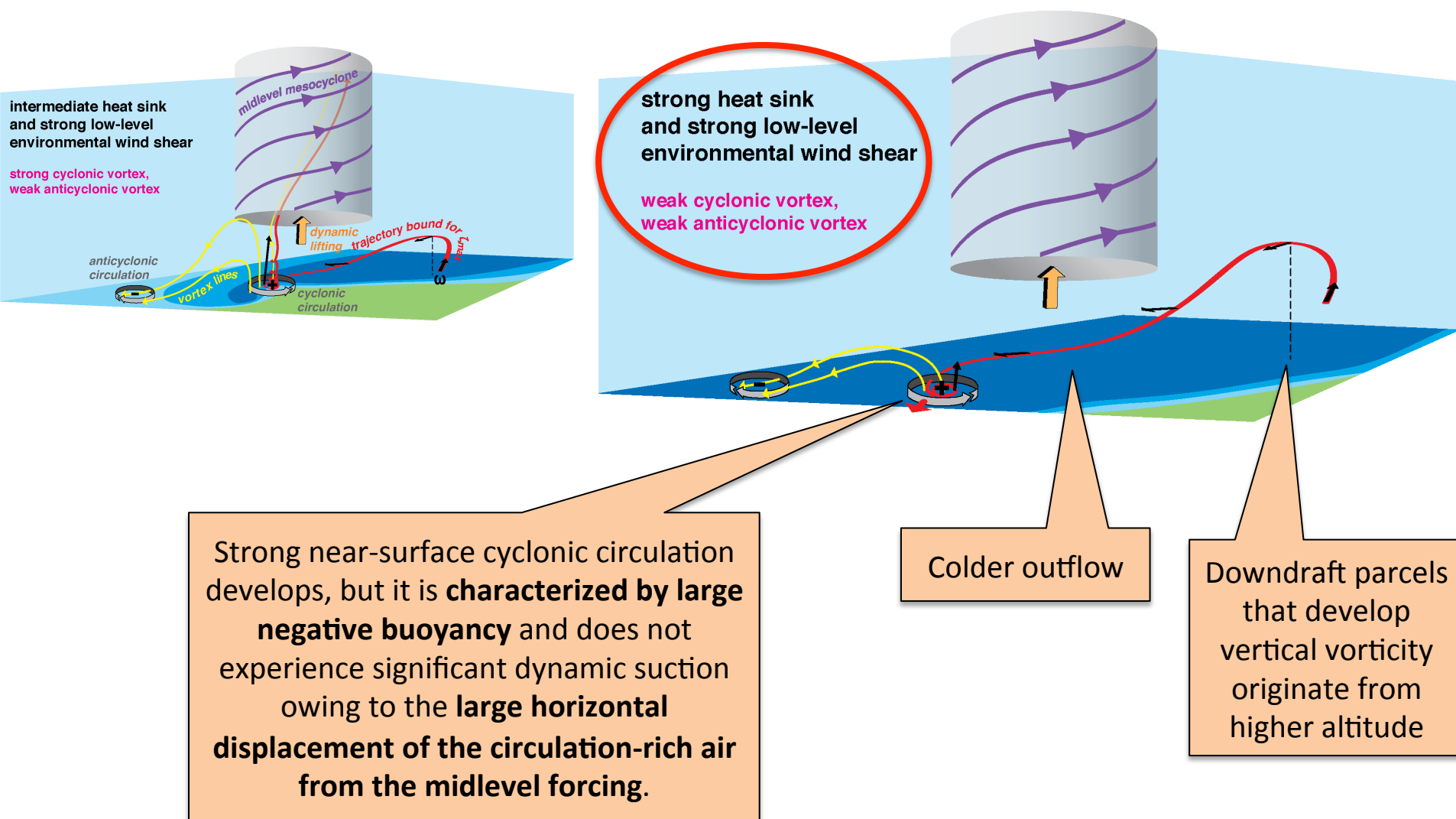
Summary

intermediate heat sink
and strong low-level
environmental wind shear

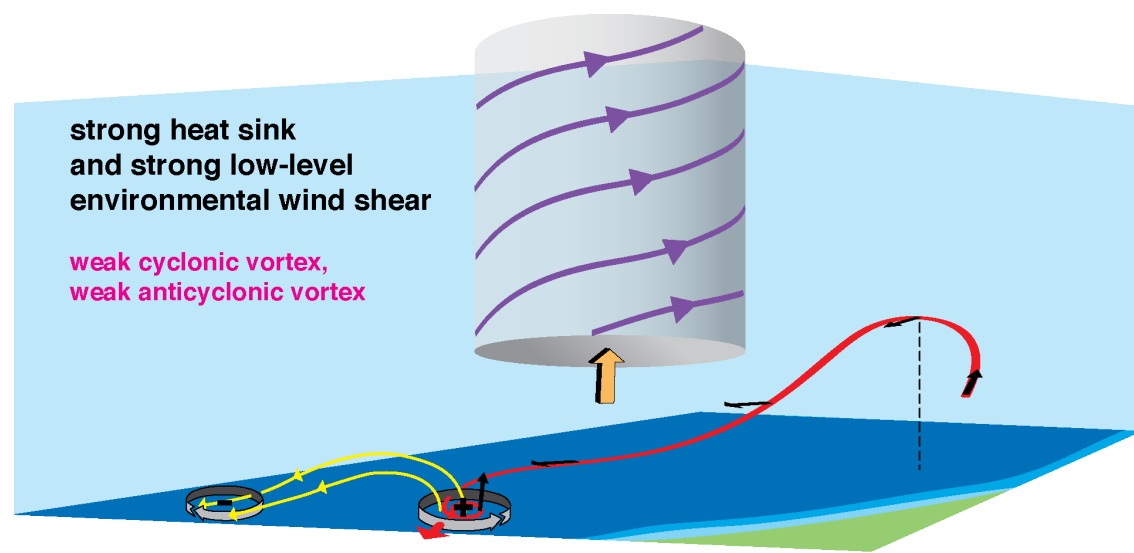
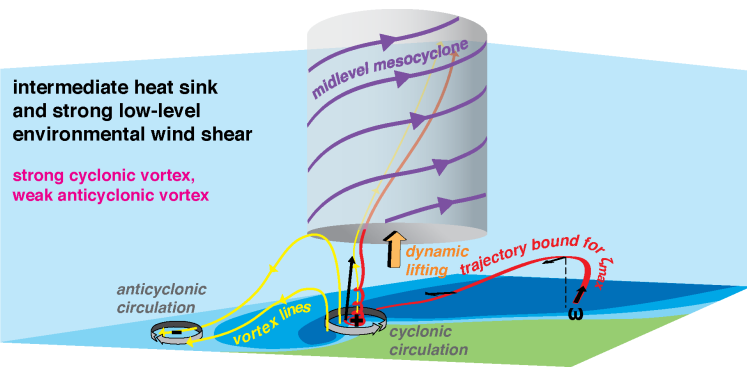
strong cyclonic vortex,
weak anticyclonic vortex



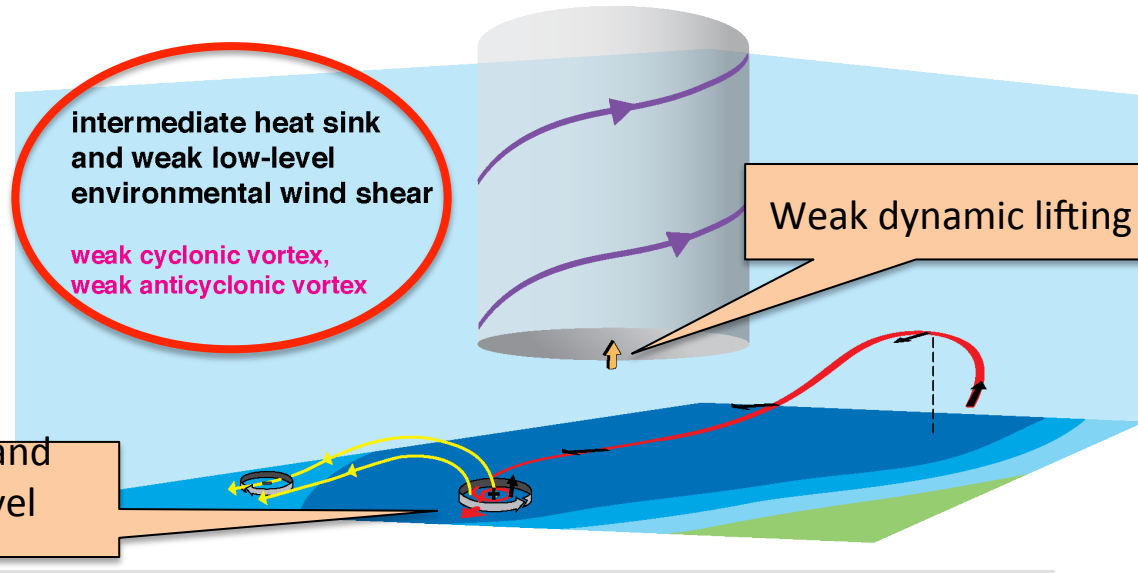
Take-home message: The development of an intense cyclonic vortex is the result of **circulation-rich, near-surface air being associated with weak negative buoyancy** and also experiencing **a large upward-directed VPPGF** owing to its proximity to the midlevel mesocyclone.



Take-home message: The development of an intense cyclonic vortex is the result of **circulation-rich, near-surface air being associated with weak negative buoyancy** and also experiencing **a large upward-directed VPPGF** owing to its proximity to the midlevel mesocyclone.



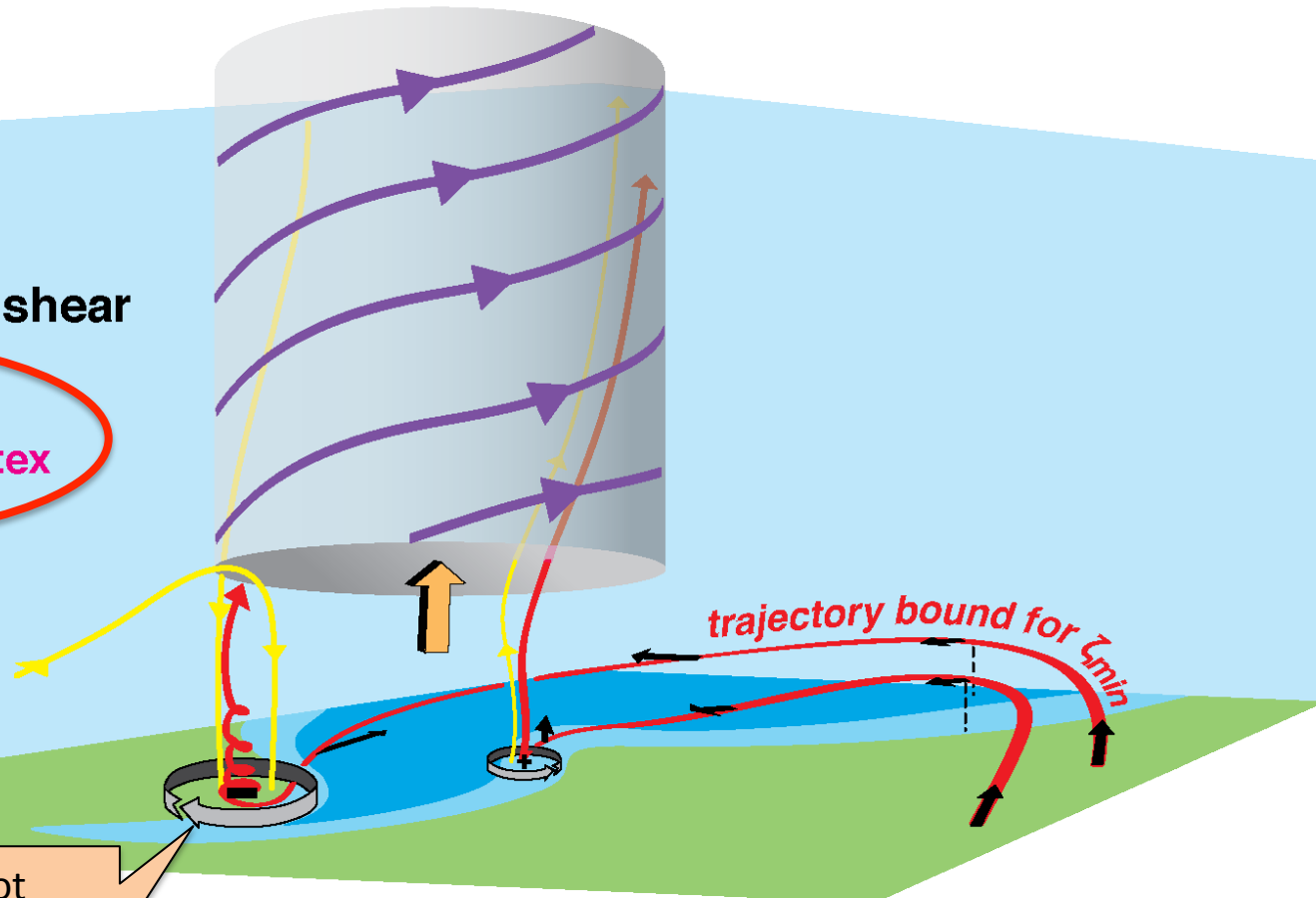
Outcome is similar to simulation with strong heat sink and strong low-level shear



Take-home message: The development of an intense cyclonic vortex is the result of **circulation-rich, near-surface air being associated with weak negative buoyancy** and also experiencing **a large upward-directed VPPGF** owing to its proximity to the midlevel mesocyclone.

weak heat sink
and strong low-level
environmental wind shear

weak cyclonic vortex,
strong anticyclonic vortex



Anticyclonic circulation is not displaced as far from the midlevel forcing as in the other simulations in which the cold pools are stronger.

Take-home message: The development of an intense cyclonic vortex is the result of **circulation-rich, near-surface air being associated with weak negative buoyancy** and also experiencing **a large upward-directed VPPGF** owing to its proximity to the midlevel mesocyclone.

Motivating Questions

- What is the role of environmental horizontal vorticity in tornadogenesis?

We are unable to refute the hypothesis that the environmental low-level horizontal vorticity's role is to promote dynamic lifting and convergence of circulation-rich, near-surface air, which is typically negatively buoyant.

- Why is the combination of strong vertical shear and high boundary-layer RH so favorable for tornadoes?

As the environmental low-level shear increases, dynamic lifting is enhanced. As boundary-layer RH increases, cold pools tend to become less negatively buoyant. The combination of a strong upward-directed VPPGF and weak negative buoyancy favors large upward accelerations and vertical vorticity stretching.

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<http://www.meteo.psu.edu/~pmm116/toysims.pdf>

